Baltic Sea Surface Temperature Analysis 2022: A Study of Marine Heatwaves and Overall High Seasonal Temperatures

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9 Abstract. In 2022, large parts of the Baltic Sea surface experienced the third-warmest to the warmest temperatures over the 10 summer and autumn temperaturesmonths since 1997. Warm temperature anomalies are a precondition forcan lead to marine heatwaves (MHWs), which are discrete periods of anomalous high temperatures relative to the usual local conditions. Here, 11 12 we describe the overall sea surface temperature (SST) conditions observed in the Baltic Sea in 2022 and provide a spatio-13 temporal description of surface MHW events based on remote sensing, model reanalyses and in-situ station data. Most The 14 most MHWs, four and locally even-up to seven MHW events, were detected in the western Baltic Sea and the Inner Danish 15 Straits, where maximum MHW intensities reached values of up to 4.6 °C above the climatological mean. The Northern 16 northern Baltic Proper and the Gulf of Bothnia were impacted by mainly by two MHWs at maximum intensities of 7.3 °C and 17 9.6 °C, respectively. Our results also reveal that MHWs in the upper layer occur at a different period than at the bottom layers 18 and are likely driven by different mechanisms. Results Model data from the case studies at two exemplary stations, 'Lighthouse 19 Kiel (LT Kiel)' and 'Northern Baltic', show a significant increase of MHW occurrences, of +0.773 MHW events per decade 20 at LT Kiel from 1989 to 2022-and of +0.64 MHW events per decade from 1993 to 2022-at Northern Baltic-, between 1993 and 21 2022. Moreover, we discuss the expected future increase of MHW occurrences is discussed increased occurrence of MHWs 22 based on a statistical analysis at both locations.

23 **1 Introduction**

Global warming has manifested itself through the<u>led to an</u> increase of ocean heat content (OHC) by about 350 ZJ in the upper 25 2000 meters from 1958 to 2019, with the year 2022 being the warmest on record <u>as of this writing</u> (Cheng et al., 2022; WMO, 26 2023). Simultaneously, marine heatwaves (MHWs), the extreme events of high water temperature (Hobday et al., 2016), have 27 increased in their frequency, duration, spatial extent and intensity during the past four decades (Sun et al., 2023). In 2022, 28 MHWs were recorded on 58 % of the ocean surface (WMO, 2023).

- The global mean air temperature in 2022 was 1.15 [1.02 1.28] °C above the 1850 1900 average, resulting in the year 2022 to be the fifth or sixth warmest year in the 173 year instrumental record (WMO, 2023). We could hence expect the marginal seas like the Baltic Sea to experience high sea surface temperatures (SST) as well. This hypothesis is supported by the study which showed that the extremely warm weather conditions in winter 2019/20 resulted in an unusually high heat content anomaly in the upper 50 m layer of the Baltic Sea (Raudsepp et al., 2022). Thus, a large scale weather pattern could have a detectable response to the physical conditions of the Baltic Sea.
- 35 Belkin (2009) showed that the Baltic Sea is one of the seas which have warmed marine ecosystems with the fastest in the last 36 decades with a recorded warming of surface temperatures of 1.35 K°C between 1982 and 2006, i.e., 0.54 °C per decade. In our 37 BSH (Belkin, 2009). SST data (operationally produced by the German Federal Maritime and Hydrographic Agency (in the 38 following BSH, product ref. no. 1 in Table 1) we see show a warming trend of 0.58 °C per decade for the period 1990–2022 39 (based on annual means derived from satellite data, not shown). Local air sea heat exchange is the main physical factor that 40 determines the surface layer water temperature and heat content in the Baltic Sea (Raudsepp et al., 2022). A permanent 41 halocline at the depth of 60-80 m (Väli et al., 2013) limits heat exchange between surface and lower layers. Further on, limited 42 water exchange of the Baltic Sea with the open ocean through the Skagerrak potentially reduces heat transport between the sea
- 43 and the ocean.
- 44 MHWs are usually detected using remotely sensed SST. High SST couldSSTs can affect phytoplankton production, while 45 unprecedented high temperatures in the subsurface layers of the sea could have even more devastating effects on the marine 46 ecosystem (Kauppi et al., 2023). Subsurface layer MHWs can be detected using temperature moorings, though unfortunately 47 these moorings provide only point measurements. Numerical models enhance the possibility for studying dynamics of MHWs 48 including their extension to the subsurface layers of the ocean. In this study we utilize remote sensing, model reanalyses and 49 in situ station data for the spatio temporal description of the MHWs in the Baltic Sea in 2022. So far, there are generally only 50 a few studies about MHWs in the Baltic Sea (Goebeler at al., 2023). Conditions that facilitate the fast warming of the Baltic 51 Sea are the limited exchange between surface and deeper layers due to a permanent halocline at a depth of 60-80 m (Väli et 52 al., 2013) and the limited water exchange between the Baltic Sea and the open ocean through the narrow Skagerrak. That is 53 why local air-sea heat exchange is the main physical factor for the surface layer water temperature and heat content in the 54 Baltic Sea (Raudsepp et al., 2022). 2022; She et al., 2020). In addition to the detailed description of MHWs in 2022 we extended 55 our study by providing climatology of the MHWs based on mooring station data. 56 Maps Global mean air temperature in 2022 was among the six warmest in the 173-year instrumental record (WMO, 2023). For 57 Europe especially, the Copernicus Climate Change Service/ECMWF (2022a) states that the air temperatures in August 2022 58 were higher than the 1991–2020 average across most of the continent, especially in a band in Eastern Europe stretching from
- 59 the Barents and Kara seas to the Caucasus. In November 2022, air temperatures were higher than the 1991–2020 average,
- 60 especially over the west, south-east and far north of Europe, and were unusually mild over the northern European seas
- 61 (Copernicus Climate Change Service/ECMWF, 2022b). These large-scale weather patterns likely lead to high sea surface

- 62 temperatures (SST) in marginal seas like the Baltic Sea and are a likely driver of MHWs. This hypothesis is further supported 63 by a study by Holbrook et al. (2019), which found that MHWs at middle and high latitude regions were driven by large-scale 64 atmospheric pressure anomalies which cause anomalous ocean warming. Stalled atmospheric high-pressure systems coincide 65 with clear skies, warm air, and reduced wind speeds. These conditions then lead to quick warming of the upper ocean and 66 increased thermal stratification due to reduced vertical mixing. 67 So far, there generally have been only a few studies on MHWs in the Baltic Sea (Goebeler at al., 2022; She et al., 2020). 68 satelliteIn this study, we show that remote sensing data revealed several SST anomalies over the entire Baltic Sea in 2022. We thus use model reanalysis and in-situ station data to provide a spatio-temporal description of the corresponding MHWs. Both 69 70 datasets contain data collected over a long enough period to also provide its own respective climatology, thereby enabling a 71 consistent representation of MHWs. While the in-situ data provides accurate point-wise measurements of the temperature at 72 selected locations, the model reanalysis data allows for a widespread analysis of MHWs over the entire Baltic Sea, including 73 their extension into subsurface layers. Furthermore, we extend our study by providing a climatology of MHWs at two specific 74 mooring locations, namely at the Lighthouse Kiel (LT Kiel) and Northern Baltic stations. The overall aim of this study is to 75 highlight the areas of the Baltic Sea that were (most) affected by MHWs and determine whether surface MHWs can propagate 76 into deeper layers and thus potentially threaten the subsurface ecosystem. Furthermore, analyzing the climatology of MHWs
- 77 can provide insight into whether the global increase in MHWs can also be expected to occur on a local scale for the Baltic Sea.

78 2 Data and Methods

79 2.1 Satellite data

80 The satellite data service at the BSH compiles daily maps of SST data (product ref. no. 1 in Table 1). -as, These have contributed 81 for example, used to studies by T the BACC Author Team (2008) and Gröger et al. (2022), are compiled daily by the satellite 82 data service at BSH (product ref. (2022). The SST no. 1-in Table 1). These data are recorded as radiances by the third generation 83 of the Advanced Very High Resolution Radiometer (AVHRR/3) in channels 4 and 5 centered at wavelengths 10.8 um and 84 12.0 um in the two thermal infrared respectively (EUMETSAT, 2015). The AVHRR/3 instruments used are flown channels 85 aboard the polar orbiting NOAA-19 and MetOp B satellites, providing a spatial resolution of 1.1-km-directly below the 86 instrument (in nadir), swath widths of 1,447-km and orbital periods of 100 minutes (EUMETSAT, 2015; Minnett et al., 2019). 87 This results in generally 8 The raw data of eight or 9 nine daytime passes over the Baltic and North Sea area per day. The raw 88 data (level 0) are received directly from EUMETSAT 90 minutes after flyover and processed using automated, standardized 89 correction procedures (atmospheric correction, cloud masking, georeferencing etc.). Additionally, each flyover is corrected 90 manually in order to preserve as much data as possible whilst eliminating any faulty or cloudy pixels. All available single 91 images from a calendar day are combined and averaged, on a single pixel basis, into one daily-mean image. These daily images 92 are then used to produce a weekly analysis on an operational basis. This weekly analysis is produced on an equidistant grid of 93 58 by 74 cells in 20 km resolution, making use of an oblique Lambert projection with an origin at 56° N, 4° W (center of the 94 BSH North Sea SST analysis). All daily satellite pixels falling into a 20x20 km grid box are considered equally when 95 calculating a weekly average for this particular 20x20 km grid box. Eventually, the result for the weekly average is smoothed 96 using a binomial filter (50 % weight for the grid box itself, 50 % for its respective neighbors). While the BSH has been carrying 97 out the processing of the satellite data itself on the 1.1 km grid has been carried out at BSH-since 1990, the operational SST 98 analysis for the Baltic Sea did not start until the autumn of 1996. The analysis of the BSH SST dataset presented in this chapter 99 is therefore constrained limited to the period from 1997–2022.

100 2.2 Station data

In-situ temperature time series from mooring stations which are located in the Baltic Sea are obtained from product ref. used for 1) model validation and 2) cross validation of the MHW computation from model data. Except no. 2 in Table 1, except for SST data from Northern Baltic (K. Hedi, FMI, pers. communication), the station data are obtained from product ref. no.)-2 in Table 1. Each available dataset has already been quality controlled by the regional production units (In Situ TAC partners, 2022). The temporal resolution varies from hourly data at the German stations and to half-hourly at the stations in the northern Baltic Proper and Gulf of Finland. Due to failures, maintenance and other circumstances, no mooring station entirely covers the period from 1st Jan 1993 until now-entirely.

108 FromOf all available mooring stations, those are we selected those which contain data from 2022 and from at least ten additional 109 years from 1993 until 2021 in-at least one depth. Out of the-then remaining seven mooring stations that contained surface 110 temperature data, two mooring stations were chosen for the additional model cross validation and the analysis of MHWs: 111 Lighthouse Kiel (LT Kiel) and Northern Baltic- (Fig. 1). Of the observation data, LT Kiel has the largest greatest time coverage 112 of observation data (rather continuously from (1989 until the present). It, missing data: 9.1 % of days). This mooring station 113 lies in the far western part of the southern Baltic, and the water depth at this mooring station there is about 12 m. On the other 114 hand, The station Northern Baltic is located in the northern Baltic Proper-and provides measurements down to water depths of 115 103.8 m. The SST observations there cover the period from 1997 until now, rendering it an ideal candidate for evaluating the 116 effects of MHWs into- (missing data: 8.0 % of days). No mooring station provides a time series in deeper layers long or 117 consistent enough to analyze subsurface MHWs, thus reducing the scope of measurement-based analysis of MHWs to the 118 surface layers.

119 2.3 Baltic Sea Physics Reanalysis Data

120 The Baltic Sea physics reanalysis product (product ref. no. 3 in Table 1) is a model dataset produced by usingbased on the

121 <u>ocean model</u> NEMO v4.0, a state of the art oceanographic modeling framework (Gurvan et al., 2019). The model system

assimilates satellite observations of SST (EU Copernicus Marine Service Product, 2022b) and in-situ temperature and salinity profile observations from the ICES database (ICES Bottle and low-resolution CTD dataset, 2022). The product provides gridded information on SST and subsurface temperature conditions. The spatial coverage is 1 nm-andnautical mile, i.e., approximately 1.8 km. The grid covers the entire Baltic Sea, including the transition zone to the North Sea, with a vertical resolution of 56 non-equidistant depth levels. This multi-year product (MYP) covers the reference period from 1993 up to 2022. The model setup is described in the Product User Manual (PUM, Ringgaard et al., 2023).

128 2.4 Model validation

129 Although the Baltic Sea physics reanalysis product (product ref. no. 3 in Table 1) has already been extensively validated in the 130 corresponding Quality Information Document (QuID; Panteleit et al., 2023), the model data is additionally validated for this 131 chapter at different mooring stations, in particular LT Kiel and Northern Baltic, using the available station data (product ref. 132 no. 2 in Table 1) and the full reference period from 1993 to 2022. Hence, in this case, only mooring stations were validated 133 which are mostly located in the southern Baltic, whereas the validation in the QuID also considered ICES data which covered 134 the entire Baltic Sea. Our evaluation confirms the results from the QuiD, for example, that the MYP underestimates the 135 temperature at the surface which, on average, results in a negative bias. Nevertheless, as can be clearly seen in Fig. 1, the 136 temperature curves at Northern Baltic and LT Kiel, respectively, show the same progression. The generally lower temperature 137 in the MYP results in a slightly lower temperature climatology and threshold (here, the 90th percentile), respectively, on which 138 the MHW detection is based. In general though, the MHWs and their respective intensities and lengths are detected equally in 139 both the station and model data.

140 While the bias at the surface is between 0.46 °C and 0.2 °C, the bias at deeper levels shifts towards more positive values. At 141 the deepest levels where observational data is available, the bias in our validation ranges from 0.22 °C to 0.43 °C which also 142 corresponds to the results in the QuID (0.36 °C to 0.26 °C). This means that the model bias varies at depth and with that 143 maybe also the accuracy. To ensure the quality of the used station data from Northern Baltic, the validation of the lower level 144 at the station BMPH2 from the QuID can be used as an approximation. BMPH2 is located only 43 km from Northern Baltic 145 and reaches a depth of 150 m (in contrast to Northern Baltic with 105 m). It is a monitoring station which is operated by vessels 146 and CTD casts are used to obtain measurements. These casts provide measurements at multiple depths, but are unevenly 147 distributed over time (Panteleit et al., 2023). Unfortunately, there is no observation data available after 2019. With a bias of 148 0.15 °C and a centered pattern root mean square difference (cRMSD) of 0.17 °C the model data fits quite well to the station 149 data (Fig. 1).

> 10 5

150 **2.5-Heat wave detection**

- Marine heatwaves (MHWs)-refer to a discrete period of unusually high seawater temperatures. While there are several definitions todescribe MHWs quantitatively describe MHWs, the most commonly used method defines them as periods when
- temperatures exceed the 90th percentile of the local climatology for five days or more (Hobday et al. 2016). Here, we apply
- 154 the python package for MHW detection and statistics by We use open-source tools to detect MHWs (Oliver-(, 2016) to the
- 155 observational data and the Matlab package by; Zhao and Marin-(, 2019) for the MHW statistics based on the in station and
- 156 model data. These packages have been shown to produce identical results (Zhao and Marin, 2019). The classification
- 157 of identified MHWs follows can be classified following Hobday et al. (2018), where in which the resulting MHW category is
- based on the maximum intensity in multiples of threshold exceedances, i.e., the local difference between the 90th percentile
- threshold and the climatology. If the threshold is exceeded by less than 2 times this local difference, the MHW is classified as moderate (Category I), at 2 to 3 times it is classified as strong (Category II), at 3 to 4 times it is classified as severe (Category
- moderate (Category I), at 2 to 3 times it is class<u>ified</u> as strong (Category II), at 3 to 4 times i
 III), and at 4 times or more times it is classified as extreme (Category IV).
- For<u>Here</u>, the assessment<u>occurrence</u> of MHWs in the Baltic domain during <u>Sea in</u> 2022 (Sect. 3.2), the climatological data of 163 1993 to 2022 as well as the data of 2022 are collected from is analyzed based on the Baltic Sea MYP (product ref. no 3 in Table 164 1). The following statistical metrics of MHWs are computed at <u>each availableevery third surface</u> grid point of the 165 reanalysisMYP, resulting in a resolution of approximately 5.4 km: cumulative intensity, mean intensity, duration of the longest 166 heatwave, number of heatwaves (frequency), maximum intensity and total days of MHW conditions.
- 167 In Then, in order to evaluate the development of those MHW metrics over time, block averages (using a block length of one
- 168 year) for each MHW metric are computed for the time series data at the two mooring stations, for both the observations
- 169 (product ref. no 2 in Table 1) and the model data. We then also compute the linear trend (95 (product ref. no 3 in Table 1) at
- 170 <u>two stations: Lighthouse Kiel and Northern Baltic. The yearly MHW metrics from observations and the model are correlated</u>
- 171 <u>for evaluation, and linear trends (95</u>% significance) <u>are calculated</u> for each of those annual MHW metrics. Finally, the
- correlation of the annual MHW metrics to the annual mean temperature based on model data was assessed using a linear least squares regression and a two-sided t-test for significance.
- All MHW assessments in the following sections use the period from 1993 to 2021 for the climatology, except for Sect. 3.2.1,
- in which the comparison of the multi-year evolution of MHWs at Northern Baltic uses the overlapping period from 1997 to
- 176 <u>2021 due to the lack of observations at this station before 1997.</u>

177 2.5 Model validation

- 178 The MYP data has already been extensively validated in the corresponding Quality Information Document (QuID; Panteleit
- et al., 2023). In this document, the MYP data are validated within the time period from 1st January 1993 to 31st December
- 180 2018. The validation shows a negative bias at the surface with a shift towards more positive values at deeper levels. A variation

- 181 of statistical values with depth is also clearly visible in the estimated accuracy number (EAN), which represents the root-mean
- 182 square difference (RMSD) of a specific depth layer. The RMSD varies between 0.29 °C at 200–400 m over 0.63 °C at the
- 183 <u>surface to 1.3 °C at 5–30 m depth.</u>
- 184 We additionally evaluated the model data in more detail using a clustering approach, which offers insights into the overall
- 185 accuracy of the model by grouping the errors. This clustering procedure employs the K-means algorithm (Raudsepp and
- 186 Maljutenko, 2022). In this evaluation, all available data within the model's domain and simulation period are considered. A
- 187 two-dimensional error space (dS, dT) is established using simultaneously measured temperature and salinity values as the
- 188 <u>foundation for clustering. Here, $dS = (S_{mod} S_{obs})$ and $dT = (T_{mod} T_{obs})$ represent the differences between the model (S_{mod} and T_{mod})</u>
- 189 and observed (S_{obs} and T_{obs}) salinity and temperature, respectively. The dataset employed in this validation study was sourced
- from the EMODNET dataset compiled by SMHI (product ref. no. 4 in Table 1). It consists of a total of 3,094,089 observations
 aligning with the simulation period of the Baltic Sea physics reanalysis (product ref. no. 3 in Table 1) and covering the years
- 192 1993 to 2022. A comprehensive explanation of the k-means-method and detailed results describing the model's accuracy can
- be found in Appendix A1. The results can be summarized as in that approximately 82 % of all validation points exhibit
- 194 relatively low temperature bias, STD, and RMSD (Table A1). The surface layer validation shows that less than 10 % of
- 195 comparison points have significant temperature errors (Figure A1c). Due to the low proportion of these validation points we
- 196 do not expect a significant impact on the determination of the surface MHWs and their statistics. Below the surface layer, i.e.,
- 197 at depths ranging from 0.5–40 m, up to 25 % of the points correspond to clusters with temperature errors greater than +/-
- 198 2.0 °C; in deeper layers, this percentage gets smaller again (Figure A1c). Consequently, we anticipate that the model reanalysis
- 199 data provide sufficiently accurate information for calculating subsurface MHWs and their statistics for the Baltic Sea as well.
- 200 The model is also validated in terms of how accurately it reproduces the MHWs of 2022 and how well it represents their
- characteristics during the overlapping time periods of data availability at the two locations (1993–2022 for LT Kiel, and 1997–
 202 for Northern Baltic). For this, the model data was compared to the available station data (product ref. no 2 in Table 1 for
 LT Kiel and K. Hedi, FMI, pers. communication for Northern Baltic) at these locations. Table 2 shows the Pearson correlation
- 204 <u>coefficients for the MHW metrics in Fig 4 between observational and model data for the two stations, which show overall</u> 205 good agreement between the model and the observation data with respect to MHW detection.
- 206 We also compared the annual temperature curves resulting from both the model and the station data at each location (Fig. A2).
- 207 Overall, the curves show the same progression. The temperature from the MYP is generally slightly lower, and consequently
- 208 this results in a slightly lower temperature climatology and threshold (here, the 90th percentile) on which MHW detection is
- based. In general, though, the MHWs and their respective intensities and lengths are detected equally in both the station and
- 210 <u>model data.</u>

211 **3 Results**

212 **3.1 Sea surface temperature anomalies in satellite data**

213 DuringIn the summer months of 2022, large parts of the Baltic Sea featured strong warm anomalies based on the BSH SST 214 analysis (product ref. no. with-1 in Table 1, Fig. 2). The highest values of were up to 3 °C above the long-term mean (1997– 215 2021) in the Bothnian Sea in June and in the Bothnian Bay in July, respectively (Fig. 2). In August however, these areas 216 exhibit rather were neutral toor exhibited cold anomalies, while the Baltic Proper as well as the Gulf of Finland and the Gulf 217 of Riga showed the warmest anomalies of +1.5 °C to 2.5 °C. InAt the beginning of autumn, the Baltic Sea is marked by a substantial east-to-west gradient regarding of SST anomalies. This is due to a series of upwelling events along theits eastern 218 219 coastlines of the Baltic Sea, shores. In November, again the whole Baltic Sea features strong warm anomalies, again with peak 220 values above +2 °C around Southern Sweden.

221 To provide some climatological context offor the observed SST anomalies in a straightforward way, we also present maps 222 of ranking the SST anomaly ranks anomalies for the summer and autumn months in of 2022 when compared to their 223 respective against the same months in previous years (right two columns of Fig. 2). It is obvious that the warm anomalies found 224 2). These anomaly rankings provide information on how extreme an anomaly of a given magnitude is. For every grid point 225 and for each calendar month, the monthly anomalies are ranked by magnitude. The warm anomalies over large parts of the 226 Baltic Sea during the summer and autumn of 2022 belong to are among the warmest eight on record for the respective months. 227 Except for In September when the coastal upwelling along the eastern coastlines led to cold anomalies contrasting the warm 228 anomalies in the western half of the Baltic Sea, it is found for basically all along the eastern shores, but the other five months 229 of the summer and fall of 2022 (June through, July and August, as well as October and November) that show large areas of the 230 Baltic Sea featured with warm anomalies belonging to that are among the highest four most pronounced on record. For In August 231 and November, we even findsee several larger areas along the coastlines of the Baltic countries as well as off the Polish coast 232 and around Gotland that featured surface temperatures being the highest ever-according to the BSH SST analysis (product ref. 233 dataset featured highest-ever surface temperatures. no. 1 in Table 1).

234 **3.2 Marine heatwaves**

Temperature MHWs describe exceptionally warm temperature anomalies are a precondition for MHWs. The. As the monthly
 overview in Fig. 2 already provides an indication of possible MHW conditions. To begin with, these are assessed by using the
 statistical_in 2022, the MHW metrics defined by Hobday et al. (2016) at each grid point based on the SST data extracted
 from are assessed using the Baltic Sea MYP (product ref. no. 3 in Table 1). Each region of the Baltic Sea experienced different
 MHW characteristics during 2022 (Fig. 3).

- 240 The most MHWs during 2022 were detected occurred in the Inner Danish Straits and the Western Baltic (Fig. 3d). Mainly 4):
- 241 <u>mainly, four</u> to <u>5five</u> MHWs were detected. <u>At, with</u> some assessed <u>grid points, locations experiencing</u> up to <u>7seven</u> MHWs

242 are detected, which leads to and a maximum of 9694 total days of MHW conditions in that area (Fig. 3f). The mean and 243 maximum intensityies of all heatwayes/MHWs in these areas the Western Baltic reached up to 3.8 °C and 4.6 °C, respectively 244 (Fig. 3b and 3e). The highest values in-mean and maximum intensity of MHWsvalues were reached in the Nnorthern Baltic 245 Proper (up to 5.3 °C, resp. 7.3 °C) as well as the Gulf of Bothnia (up to 6.5 °C, resp. 9.6 °C) and in the Bothnian Sea and 246 Bothnian Bay (Fig. 3b and 3e), though these regions were affected mainly by only two MHWs. The maximum intensities were 247 evenintensity in the Bothnian Bay even reached 9.6 °C, the highest within the entire studied period from 1993 to 2022. Those 248 regions were impacted by mainly two MHWs, but at some individual grid points up to six MHWs are identified with a 249 maximum value of up to 63 days with MHW qualification. While the duration of the The longest MHW is similar in both 250 regions (Fig. found in the Baltic Proper (32 days), followed by the Bothnian Sea (31 days) and the Inner Danish Straits (29 251 days) (Fig. 3c)-, tThe highest values of cumulative intensity (of a single MHW) are found in the Gulf of Bothnia in the Kvarken 252 (), with up to 119.3 days °C), while the values in the Western Baltic and Inner Danish Straits are lower (up to 64 days °C) (Fig. 253 , are found in the Kvarken, a strait between the Bothnian Sea and 3a).

Compared with Fig. 2, the Bothnian Bay (Fig. 3a). MHW with the highest cumulative intensity in the Gulf of Bothnia derives
 from the temperature anomaly in November. The numerous MHWs in the Western Baltic and the large number of days with
 MHW conditions connected with them occurred in both the summer and autumn months.

257 **3.2.1 Multi-year evaluation of MHW metrics**

258 Next, we assess the frequency and other characteristics of the MHWs that occurred in 2022 in a climatological context based 259 on both observations and model data for the two stations, LT Kiel (forbased on the overlapping climatology period 1993-260 20221, Fig. 4a–h) and Northern Baltic (forbased on the overlapping climatology period 1997–20221, Fig. 4 i–p). Overall, the 261 results for the yearly MHW metric calculation are well correlated between the observations and the model data (Table 2). 262 In 2022, in total of five MHWs (four in the MYP) occurred throughout the year at LT Kiel, spread out through the year 263 (Fig. 1aA2a). Though none of the 2022 MHWs them was extraordinarily long or intense at LT Kiel, the station data time series 264 of yearly MHW metrics shows that, based on observational data, the number of MHW occurrences in 2022 was the second 265 highest after 2020there since 1989 (Fig. 4a). The time series of MHW frequencyies per year suggests that the occurrence of 266 MHW events has increased over the last three decades (Fig. 4a). Thise trend of computed from model data is +0.773 MHWs 267 per decade becomes statistically significant when all for the available station data from 1989 period 1993-2022-is taken into 268 account. The number of MHW events per year is positively correlated (R=0.76) with the increasing annual mean SST at that is 269 mooring station (Fig. 4b). The maximum (Fig. 4c) and cumulative intensities (Fig. 4e) of observed MHWs do not show a clear 270 trend and are not correlated to risingthe warming annual mean temperatures (Fig. 4d and Fig. 4f). There is no significant trend 271 in total MHW days (Fig. 4g) at LT Kiel, but a positive correlation (R=0.71) with rising average temperatures (Fig. 4h).

272 For Northern Baltic, neither the station data nor the model data exhibits a statistically significant trend in MHW events for the 273 overlapping period (Fig. 4i). But when all of the available model data from 1993–2022 is taken into account, the trend in MHW 274 occurrences becomes significant at the 95 % level, with +0.64 MHWs per decade. Again, the number of events is positively 275 correlated with annual mean temperature (R=0.58, Fig. 4j). The highest maximum MHW intensities were recorded in recent 276 years (2016, 2018, 2021, 2022)), with 2022 showing the highest intensity of any MHW-with, at 7.3 °C (model data) to 7.4 °C 277 (station data) above the climatologically expected temperature (Fig. 4k,l, see also Fig. 1bA2b). The cumulative MHW 278 intensities show no clear trend or correlation with annual mean temperatures at this station (Fig. 4m,n). In terms of total MHW 279 days, 2018 was exceptional shows the highest numbers (Fig. 40), but otherwise no trend is detectable for this metric, though 280 there is positive correlation with annual mean temperatures (R=0.56, Fig. 4p).

281 **3.2.2 Analysis of vertical MHW distribution at Northern Baltic**

At Northern Baltic, which is more than 104<u>about 103</u> m deep and located in the Western Baltic Proper, the surface temperature has been <u>measured</u>_continuously <u>measured</u>_over several decades. Unfortunately, no <u>quality-controlled</u>_temperature measurements exist <u>infor the</u> lower layers. In Sect. 2.4 we showed that the <u>at this station</u>. The model <u>data coincidesvalidation</u> shows that, at other locations, the model represents temperatures generally well-with observational <u>data</u>, both at the surface and in the lower layers. <u>Thus, inIn</u> order to obtain further insights into the heat wave propagation towards the seafloor, we analyzed the MYP model data (product ref. no. 3 in Table 1) along the entire water column.

288 A seasonal SST signal is clearly visible in Fig.- 5a. In general, the temperature tends to decrease with depth while the bottom 289 temperature is relatively cold and uniform. In late springearly summer (June), a so-called cold intermediate layer (CIL) at a 290 depth of 20 60 m, which is), defined as a minimum of temperature between the thermocline and the perennial halocline 291 (Chubarenko et al., 2017; Dutheil et al., 2022), is formed. The upper boundary of the CIL is in good agreement with the mixed 292 layer depth. The CIL2022), develops at a depth of 20–60 m and acts as a barrier between the surface and bottom water body 293 and has bodies. At Northern Baltic, the upper boundary of the CIL coincides with the mixed layer depth (MLD), which is 294 depicted in Fig. 5b-c. Starting from around June, a layer of water with a significantly lower temperature (Fig. 5b) of 0.3 °C to 295 4.5 °C as the climatological mean. As shown in Fig. 5c and Fig. 5d, the intensity of the MHW tends to decrease as the depth 296 increases. MHWs in regions close to the seafloor were detected during specific periods from February to April, September to 297 October and in December. During July, a one day extreme MHW (Category IV) event was observed at the surface with 7.4 °C 298 above the climatological mean temperature, followed by further three days with a severe MHW (Category III). A few weeks 299 prior to this MHW (and in 10.8 m depth also afterwards), the temperature was already significantly higher than the 300 climatological mean, but (up to -7 °C deviation) is found just below the MLD (Fig. 5b), which suggests that the CIL was 301 significantly colder at this time in 2022. This also coincides with the onset of significantly higher temperatures at the surface 302 compared to the climatological mean, though these were initially not high enough to result in a MHW (Fig.- 5e-and 5f). The

303 elevated temperatures started with a significant temperature jump of 5- $^{\circ}C$ above the climatological mean, followed by an 304 abrupt and substantial indecreases and deincreases in temperature over a short period (Fig. 5e). This eventually leads to a 305 MHW which lasts for 15 days starting from the end of June and which contains a one-day extreme MHW (Category IV) event 306 at a temperature of 7.4 °C above the climatological mean, followed by a severe MHW (Category III) for another three days. 307 Significantly high temperature deviations can also be observed at a depth of 10.8 m, i.e., at the MLD, after July 2nd, just after 308 the Category IV MHW at the surface. However, these temperature deviations did not result in a MHW at this depth. Following 309 thise extreme heatwave event at the surface, a comparably weaker MHW is observed can be detected in mid-August. The 310 weaker August MHW is observable at both 0.5- m (Fig.- 5e) and 10.8- m (Fig.- 5f), while the June). Thus, this weaker MHW 311 is already no longer noticeable below the mixed layer depth, penetrates past the MLD into slightly deeper levels before reaching 312 the comparably cold layer of water underneath. 313 As shown in Fig. 5c and Fig. 5d, the intensity of the MHW tends to decrease as the depth increases. Four MHWs in regions 314 close to the seafloor (i.e., below 60 m) were detected during specific periods from February to April, September to October, 315 and in December. These MHWs are mostly moderate, with temperatures reaching up to 1.59 °C above the climatological mean. 316 Only three days at the end of September can be classified as a Category II MHW in one specific depth-layer close to the 317 seafloor. In the bottom-most depth-layer, the corresponding subsurface MHW is interrupted by five days of temperatures below 318 the 90th percentile. However, as the temperatures are only slightly below the threshold and the MHW criteria are still met in 319 the depth-layers above, one might still count this as one continuous MHW. Furthermore, Fig. 5c also shows isolated Category 320 I MHWs at depths between 20 and 50 m.

321 4 Discussion and Conclusions

322 During August and November 2022, record-warm sea surface temperatures were observed in substantial areas inof the Baltic 323 Sea proper. Large parts of the Baltic Sea exhibited the third-warmest to the warmest temperatures in summer or autumn 324 temperaturesmonths since 1997. Both periods, in August and November, coincided with atmospheric temperature anomalies. 325 In August air temperatures were higher than Over the 1991 2020 average across most of Europe, especially in Eastern Europe, 326 in a band stretching from the Barents and Kara seas to the Caucasus (Copernicus Climate Change Service/ECMWF, 2022a). 327 In November entire year of 2022, air temperatures were higher than the 1991 2020 average especially over the west, south-328 east and far north of Europe and unusually mild over the northern European seas (Copernicus Climate Change 329 Service/ECMWF, 2022b). These atmospheric surface temperature anomalies are a likely driver for MHWs since they seem to 330 coincide with the observed marine surface temperature anomalies. Holbrook et al. (2019) found that the MHWs they studied 331 at middle and high latitude regions were driven by large scale atmospheric pressure anomalies which cause anomalous ocean 332 warming. Stalled ridges of atmospheric high pressure systems coincide with clear skies, warm air, and reduced wind speeds. 333 These conditions lead to quick warming of the upper ocean and increase thermal stratification due to reduced vertical mixing.

334 In 2022, MHWs occurred in all marginal seas of Europe where the Baltic Sea was not an exception. The the distribution of 335 quantity and intensity of MHWs within the Baltic Sea is twofold: Upup to seven individual MHW occurrences were both 336 recorded and as well as simulated in the south-western part of the Baltic Sea, which lead to the fact that and as a result this 337 region experienced the maximum number of total MHW days in the of anywhere in the Baltic Sea in 2022. In the northern 338 Baltic Sea-in 2022. In, the northern part, the total number of MHWs was lower, with locally some locations registering only 339 one MHW-being detected. Remarkably; remarkably, however, this one MHW-also led to the highest mean and maximum 340 MHW intensities in the Baltic Sea since the reanalysis started in 1993. In some areas, in the Bothnian Bay, the Baltic Sea 341 MYP revealed temperatures that exceeded 9 °C above the 90th percentile of the climatologically expected temperature values-342 (Fig. 3d,e). This can be considered an extraordinarily high MHW intensity, since maximum SST anomalies above 5 °C have 343 only been observed in about 5 % of the global ocean, and MHW intensities normally peak at 2.5 °C to 3.7 °C (Sen Gupta et 344 al., 2020). In our case, the area in the Bothnian Bay experienced a short period with southerly winds and air temperatures up 345 to 28 °C at the end of June 2022 (SMHI, 2023), which led to a short, but very intense MHW in the shallow areas of the Bay. 346 At our two exemplary stations aA significant increase of MHW occurrences is detectable over time- at our two exemplary 347 stations, of +0.773 MHW events per decade at LT Kiel and +0.64 MHW events per decade at Northern Baltic. There is a 348 statistical relationship between bothBoth MHW frequency and the total number of MHW days with are statistically related to 349 rising mean temperatures. This confirms that, also in the Baltic Sea, an increasing number of MHWs can be expected in the 350 future in the Baltic Sea, too, due to global warming (Frölicher et al., 2018; Oliver et al., 2019). Adverse The adverse impact of 351 MHWs to different on the ecosystem's various trophic levels of the ecosystem is has been widely documented (Smale et al., 352 2019; IPCC, 2022; Smith et al., 2023). The Baltic Sea, which has a relatively vulnerable ecosystem, could experience a 353 significant negative impact due to from MHWs (Kauppi and Villnäs, 2022; Kauppi et al., 2023), 2023), and the analysis of 354 subsurface MHWs opens up further potential ways to study their effects. At the Northern Baltic mooring station, MHWs were 355 found at the surface, propagating into deeper layers until reaching the CIL, and some were also detected close to the seafloor. 356 Isolated MHWs were also observed at depths of between 20 and 50 m. However, these are subject to higher uncertainty 357 compared to the ones in the surface and bottom layers due to a higher uncertainty in modeling variability in the pycnocline 358 (OuID: Panteleit et al., 2023). Among the possible reasons for the development of the four MHWs close to the seafloor at 359 Northern Baltic could be vertical heat transport from the surface or a lateral transport of warmer water due to bottom currents, 360 for example. However, a more detailed evaluation would be required to assess their precise cause. Potential avenues for future studies opens up by include examining whether the August MHW in 2022 could and how surface 361 362 MHWs are able to propagate into the deeper water masses close to the halocline at Northern Baltic and also regarding as well 363 as examining the correlation between the strength (i.e., the classification category) of the MHW and its propagation into deeper 364 water masses and the strength (i.e. classification) of the MHW. Further. At Northern Baltic, severe and extreme MHWs 365 occurred at the surface when the CIL was particularly cold compared to the climatology. This therefore raises questions of

 $\frac{366}{10}$ whether a strong CIL might be linked to the development of MHWs at the surface and whether the one might even favor the $\frac{10}{10}$

 $\frac{development of the other. Additional studies could determine if also focus on the positive feedback on the bottom temperature,$ $<math display="block">\frac{development of the other. Additional studies could determine if also focus on the positive feedback on the bottom temperature,$ $<math display="block">\frac{development of the other. Additional studies could determine if also focus on the positive feedback on the bottom temperature,$ $<math display="block">\frac{development of the other. Additional studies could determine if also focus on the positive feedback on the bottom temperature,$ $<math display="block">\frac{development of the other. Additional studies could determine if also focus on the positive feedback on the bottom temperature,$ $<math display="block">\frac{development of the other. Additional studies could determine if also focus on the positive feedback on the bottom temperature,$ $<math display="block">\frac{development of the other. Additional studies could determine if also focus on the positive feedback on the bottom temperature,$ $<math display="block">\frac{development of the other. Additional studies could determine if this phenomenon can also be found in following other$ $<math display="block">\frac{development of the other. Additional studies could determine if this phenomenon can also be found in following other$ $<math display="block">\frac{development of the other. Additional studies could determine if this phenomenon can also be found in following other$ $<math display="block">\frac{development of the other. Additional studies could determine if this phenomenon can also be found in following other$ $<math display="block">\frac{development of the other. Additional studies could determine if this phenomenon can also be found in following other$ $<math display="block">\frac{development of the other. Additional studies could determine if the other studies the other studies could determine the other studies of the other studies the other studies of the other studies the other studies of the other s$

373 Appendix A1

- We apply a clustering approach to evaluate the precision of the hydrodynamic model. This technique offers insights into the overall accuracy of the model by grouping the errors. The clustering procedure employs the K-means algorithm, a type of unsupervised machine learning (Jain, 2010). The original explanation of this technique can be found in a study by Raudsepp and Maljutenko (2022). In our evaluation, all available data within the model's domain and simulation period are considered, even if the observation data is unevenly distributed or occasionally sparse. This strategy enables us to assess the model's quality at each specific location and time instance at which measurements have been acquired.
- 380 Initially, a two-dimensional error space (dS, dT) was established using simultaneously-measured temperature and salinity 381 values as the foundation for clustering. Here, $dS = (S_{mod} - S_{obs})$ and $dT = (T_{mod} - T_{obs})$ represent the differences between the model 382 (S_{mod} and T_{mod}) and observed (S_{obs} and T_{obs}) salinity and temperature, respectively. The dataset employed in this validation 383 study was sourced from the EMODNET dataset compiled by SMHI (product ref. no. 4 in Table 1). It consists of a total of 384 3,094,089 observations aligning with the simulation period of the Baltic Sea physics reanalysis (product ref. no. 3 in Table 1) 385 and covering the years 1993 to 2022. For each observation, we extracted the nearest model values from the reanalysis dataset. 386 The next stage involves choosing the number of clusters, and for simplicity we opted in advance for five clusters. Subsequently, 387 the third step entails conducting K-means clustering on the two-dimensional errors. This clustering process is applied to the 388 normalized errors achieved through separate normalization for temperature and salinity errors using the corresponding standard 389 deviations. The K-means algorithm then identifies the centroids' positions within the error space for the predetermined number 390 of clusters. These centroids' locations signify the bias of the error set for each cluster. In the fourth step, statistical metrics for 391 non-normalized clustered errors are computed. Standard deviation (STD), root mean square deviation (RMSD) and the 392 correlation coefficient are examples of common statistics that can be calculated for the parameters associated with each cluster. 393 The fifth step involves examining the spatio-temporal distributions of errors associated with different clusters. During the 394 creation of the error space, we retained the coordinates of each error point (dS, dT)(x, y), allowing us to map the errors of each 395 cluster back onto the locations where the measurements were conducted. To achieve this, the model domain is partitioned into 396 horizontal grid cells (i, i) of $27x27 \text{ km}^2$ in size. Subsequently, the number of error points attributed to various clusters at each 397 grid cell (i, j) is tallied. The total number of error points linked to the grid cell (i, j) is the sum of points from each cluster. The

- ³⁹⁸ proportion of error points in each grid cell affiliated with cluster k is determined by the ratio of the number of error points of
- 399 <u>cluster k to the total number of error points in each grid cell.</u>
- 400 401 Figure respectA1 displays the results of the K-means clustering for non-normalized errors. Table S1 presents the corresponding 402 metrics. Within cluster k=5, the salinity and temperature values closely align with the observations, with a bias of dS=-403 0.40 g/kg and dT=-0.02 °C, respectively. This cluster encompasses 57 % of all data points. The points are distributed 404 throughout the Baltic Sea and the great majority of them exceed 0.5 (Figure A1b). Clusters k=3 and k=4 exhibit relatively even 405 spatial distributions across the Baltic Sea, accounting for 11 % and 8 % of the points, respectively. These clusters are 406 particularly noteworthy due to their relatively high temperature biases and variability, both of which are crucial for the 407 calculation of marine heatwaves. The clusters k=1 and k=2 represent points with low temperature but a high salinity error 408 (Table A1). Spatially, these points are predominantly located in the southwestern Baltic Sea (Figure A1b), which points to the 409 occasional underestimation or overestimation of the inflow/outflow salinity.
 - 410 Collectively, approximately 82 % of all validation points exhibit relatively low temperature bias, STD and RMSD (Table A1).
 - 11 The surface-layer validation shows that less than 10 % of comparison points have significant temperature errors (Figure A1c).
 - 412 Due to the low proportion of these validation points, we do not expect a significant impact on the ecosystem in subsurface
 - 413 layersdetermination of surface MHWs and their statistics. Below the surface layer, i.e., at depths ranging from 0.5–40 m, up
 - to 25 % of the points correspond to clusters k=3 and k=4 (Figure A1c). Consequently, we anticipate that the model reanalysis
 - 415 <u>data provides sufficiently accurate information for calculating subsurface MHWs and their statistics for the Baltic Sea</u>.

416 Data availability

417 This study is based on public databases and the references are listed in Table 1.

418 Author contribution

419 The idea for and concept of behind this chapter was were formed by Anja Lindenthal, Claudia Hinrichs, Priidik Lagemaa, Helen 420 E. Morrison and Urmas Raudsepp. The data curation was done by Eefke M. van der Lee and Tim Kruschke for the data from 421 product ref. no. 1 in Table 1, by Claudia Hinrichs and Tabea R. Panteleit for the data from product ref. no. 2 in Table 1 and by 422 Simon Jandt-Scheelke and Tabea R. Panteleit for the data from product ref. no. 3 in Table 1. The formal analyses of the datasets 423 and the resulting investigations were performed by Anja Lindenthal, Claudia Hinrichs, Simon Jandt-Scheelke, Tim Kruschke 424 and Tabea R. Panteleit. Additional The k-means model validation was performed by Tabea R. Panteleit. Urmas Raudsepp and 425 Ilja Maljutenko. Claudia Hinrichs, Simon Jandt-Scheelke, Ilja Maljutenko, Tim Kruschke and Tabea R. Panteleit were 426 responsible for the visualization of the data. Anja Lindenthal, Claudia Hinrichs, Simon Jandt-Scheelke, Tim Kruschke, Eefke 427 M. van der Lee, Tabea R. Panteleit and Urmas Raudsepp were involved in the original draft preparation, while the. The final

428 manuscript was-finally reviewed and edited by Claudia Hinrichs, Priidik Lagemaa, Helen E. Morrison and Urmas Raudsepp

429 with contributions from all co-authors.

430 **Competing interests**

431 The authors declare that they have no conflict of interest.

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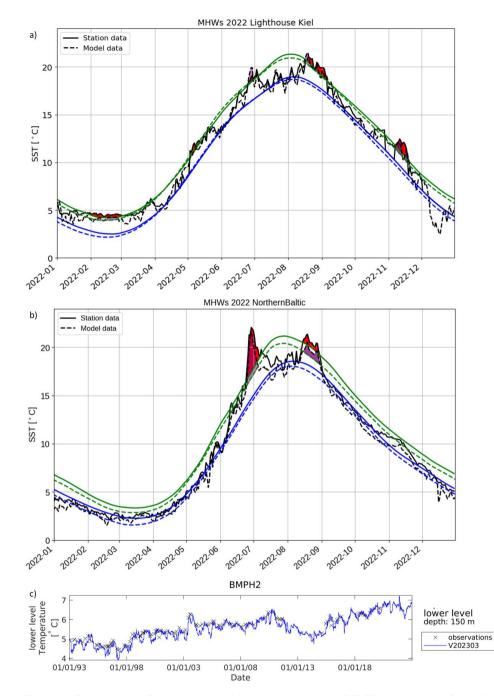
552 <u>Tables</u>

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554 Table 1: Product Table

Product ref. no.	Product ID & type	Data access	Documentation		
1	BSH Sea Surface Temperature (AVHRR/3); Satellite data	Upon request; overview and contact data via https://www.bsh.de/EN/TO PICS/Monitoring_systems/ Remote_sensing/remote_se nsing_node.html	https://www.bsh.de/DE/THEMEN/Beob achtungssysteme/Fernerkundung/ferner kundung_node.html		
2	INSITU_GLO_PHYBGCWA V_DISCRETE_MYNRT_013_ 030; In-Situ Near-Real-Time Observations	EU Copernicus Marine Service Product (2022a)	Quality Information Document (QUID): Wehde et al., 2022Wehde et al. (2022) Product User Manual (PUM): In Situ TAC partners, 2022In Situ TAC partners (2022)		
3	BALTICSEA_MULTIYEAR_ PHY_003_011; Numerical models	EU Copernicus Marine Service Product (2023)	Quality Information Document (QUID): Panteleit et al., 2023Panteleit et al. (2023) Product User Manual (PUM): Ringgaard et al., 2023Ringgaard et al. (2023)		
<u>4</u>	EMODNET CHEMISTRY B altic_Sea_aggregated_ eutrophication_and_acidity_dat asets 1902-2017_v2018; Observations	<u>SMHI (2019)</u>	Buga et al. (2018), Giorgetti et al. (2020)		

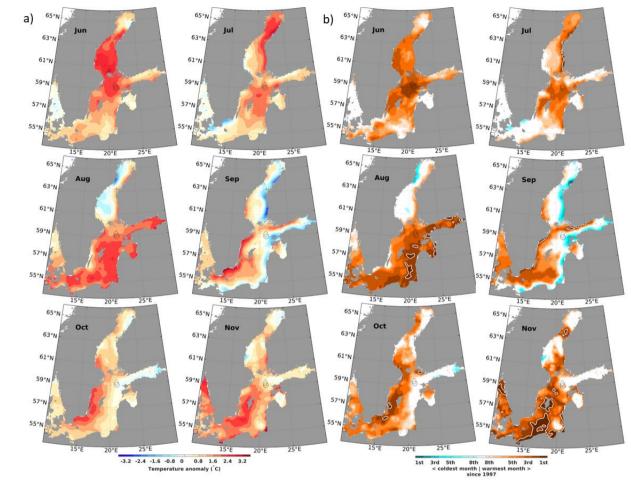
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558 Figure 1: Comparison of station data with model data at (a) LT Kiel (product ref. no.-2 and 3 in Table 1), (b) Northern Baltic (product ref. no.-2 and 3 in Table 1) and (c) BMPH2 (from Pantelcit et al., 2023). The dashed lines in (a) and (b) correspond to the 560 model (product ref. no.-3 in Table 1), while the continuous lines correspond to the station data (product ref. no. 2 in Table 1). In 561 blue, the climatological mean is shown. The green lines show the 90th percentile threshold for MHW detection and the black lines 562 are the respective 2022 temperature data. The purple (model data) and red (station data) marked areas show the detected MHWs

563 in 2022. The reference period for LT Kiel (a) is 1993-2021 and 1997-2021 for Northern Baltic (b), (c) shows the validation at the 564 station BMPH2 at a depth of 150 m. The model data is shown in blue and the measured data are displayed with the black crosses.





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Table 2: Pearson correlation coefficients from linear regression between the MHW metrics computed from the station data and the model data at the stations Lighthouse Kiel and Northern Baltic.

<u>Station</u>	<u>common</u> <u>climatology</u> <u>period</u>	MHW count	<u>MHW max</u> intensity	MHW cumulative intensity	total MHW days
Lighthouse Kiel	<u>1993-2021</u>	<u>0.82</u>	<u>0.88</u>	<u>0.66</u>	<u>0.93</u>
Northern Baltic	<u>1997-2021</u>	<u>0.74</u>	<u>0.89</u>	<u>0.82</u>	<u>0.94</u>

571 572 573 Table 3: Statistical MHW parameter values in various subregions of the Baltic Sea for 2022 based on the model data from the Baltic Sea MYP (product ref. no. 3 in Table 1) using daily values of SST between 1st January 1993 and 31st December 2022. The

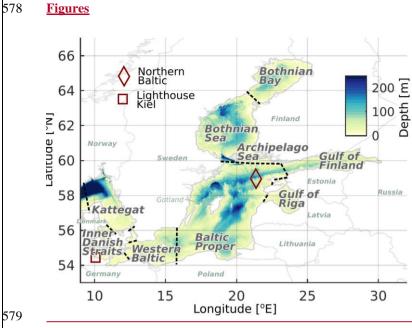
climatological period covers the years 1993 to 2021.

	<u>Kattegat</u>	<u>Inner</u> Danish <u>Straits</u>	<u>Western</u> <u>Baltic</u>	Baltic Proper	<u>Gulf of</u> <u>Riga</u>	<u>Gulf of</u> <u>Finland</u>	<u>Archipe-</u> lago Sea	<u>Bothnian</u> <u>Sea</u>	<u>Bothnian</u> <u>Bay</u>
Cumulative intensity of longest <u>MHW</u> / °C days	<u>81.5</u>	<u>63.8</u>	<u>64</u>	<u>79.4</u>	<u>63</u>	<u>66.5</u>	<u>61.1</u>	<u>119.3</u>	<u>85.1</u>
<u>Mean</u> intensity / °C	<u>3.6</u>	<u>3.5</u>	<u>3.8</u>	<u>5.3</u>	<u>4.9</u>	<u>5.8</u>	<u>4.5</u>	<u>6.4</u>	<u>6.5</u>
Duration of longest MHW / days	<u>24</u>	<u>29</u>	<u>26</u>	<u>32</u>	<u>17</u>	<u>17</u>	<u>21</u>	<u>31</u>	<u>20</u>
<u>Number of</u> <u>MHWs</u> (modal) per year	<u>1-6 (3)</u>	<u>2-7 (4)</u>	<u>2-7 (5)</u>	<u>1-7 (3)</u>	<u>1-4 (3)</u>	<u>1-4 (2)</u>	<u>2-4 (3)</u>	<u>1-6 (2)</u>	<u>1-5 (2)</u>
<u>Maximum</u> intensity / °C	<u>4.5</u>	<u>4.2</u>	<u>4.6</u>	<u>7.3</u>	<u>5.9</u>	<u>6.8</u>	<u>5.1</u>	<u>8.6</u>	<u>9.6</u>
<u>Total days of</u> <u>MHW</u> <u>conditions /</u> <u>days</u>	<u>56</u>	<u>86</u>	<u>94</u>	<u>79</u>	<u>50</u>	<u>48</u>	<u>55</u>	<u>63</u>	<u>47</u>

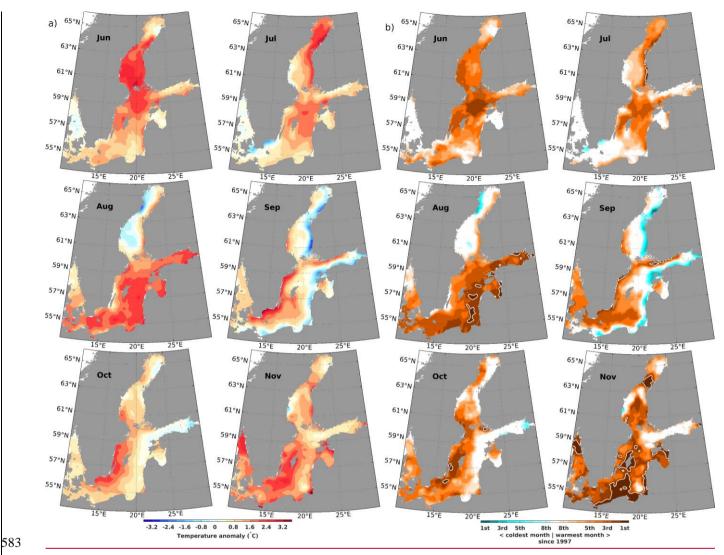
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575 Table A1: The share (%), bias, root-mean-square error (RMSE), standard deviation (SD), and correlation coefficient (Corr) for 576 each of the five clusters.

	Shares	Bias		<u>SD</u>		<u>RMSE</u>		Corr		
k	<u>%</u>	<u>dS</u>	dT	<u>dS</u>	dT	<u>S</u>	<u>T</u>	<u>S</u>	<u>T</u>	<u>dSdT</u>
		<u>(g/kg)</u>	(°C)	<u>(g/kg)</u>	<u>(°C)</u>	<u>(g/kg)</u>	<u>(°C)</u>			
1	<u>18.6</u>	<u>-4.14</u>	-0.26	<u>1.80</u>	<u>0.85</u>	<u>4.51</u>	<u>0.89</u>	<u>0.90</u>	<u>0.78</u>	<u>-0.09</u>
2	<u>7.4</u>	<u>3.53</u>	<u>0.39</u>	<u>2.16</u>	<u>1.06</u>	<u>4.14</u>	<u>1.13</u>	<u>0.93</u>	<u>0.75</u>	<u>-0.11</u>
3	<u>10.5</u>	-0.62	2.58	2.12	<u>1.28</u>	2.21	2.88	<u>0.97</u>	0.58	<u>-0.06</u>
4	<u>6.3</u>	<u>0.27</u>	<u>-2.29</u>	<u>1.97</u>	<u>1.21</u>	<u>1.99</u>	<u>2.59</u>	<u>0.95</u>	<u>0.71</u>	<u>-0.14</u>
<u>5</u>	<u>57.2</u>	<u>-0.40</u>	<u>-0.02</u>	<u>0.83</u>	<u>0.54</u>	<u>0.92</u>	<u>0.54</u>	<u>0.99</u>	<u>0.89</u>	<u>0.07</u>

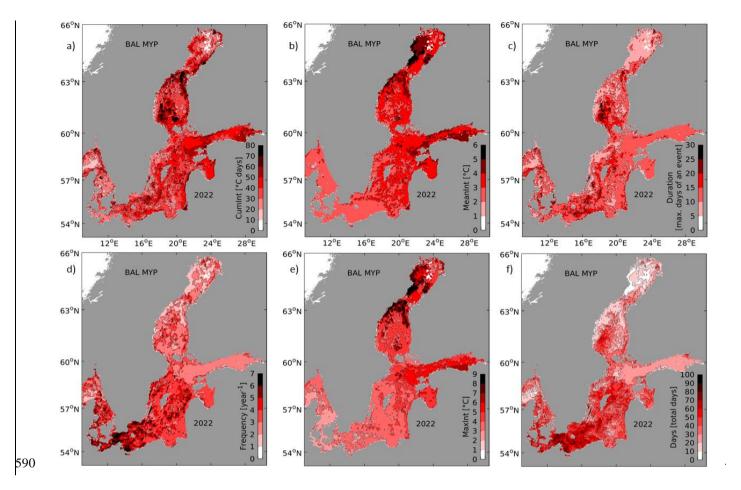


- Figure 1: Map of the Baltic Sea with relevant locations mentioned in the study. Boundaries between subregions are marked with dashed lines.

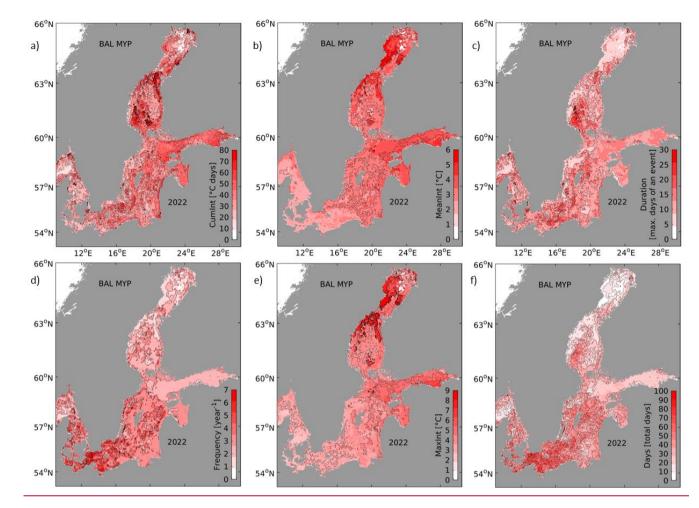


584 <u>Figure 2:</u>

Figure-2: Anomalies (difference to climatology of 1997-2021) of SST for the Baltic Sea according to the BSH SST analysis (product ref. no 1 in Table 1) during the summer and autumn months in 2022 (a) and ranks of these SST anomalies (b) when compared to the full dataset starting in 1997. In (b), <u>Bb</u>rownish (cyan) colors denote anomalies belonging to the warmest (coldest) eight anomalies found since 1997. Record warm anomalies (rank 1) are highlighted by white contours. <u>Locations of the in situ observations</u> discussed in this chapter are marked by a square (LT Kiel), a diamond (Northern Baltic), and a circle (BMPH2), respectively.



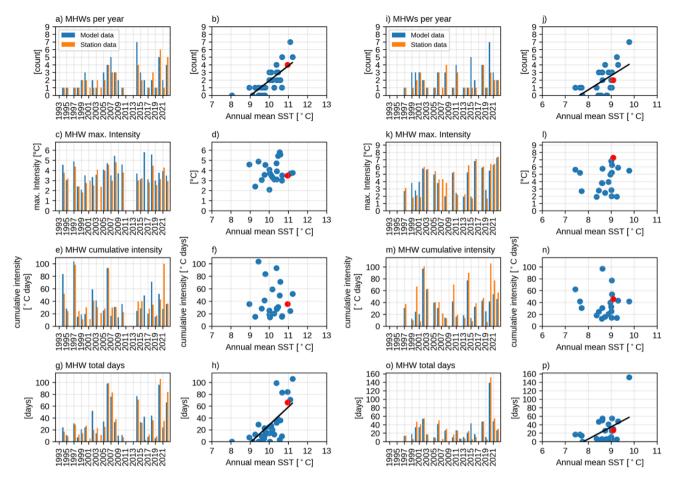
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592 Figure 13: Statistical metrics of MHWs in 2022 in the Baltic Sea based on SST data of the Baltic Sea MYP (product ref. no. 3 in 593 Table 1) with the climatological period covering the years 1993 to 20221 - (a) cumulative intensity of the longest heatwave, (b) mean 594 intensity, (c) duration of the longest heatwave, (d) number of heatwaves during 2022, (e) maximum intensity during the longest heatwave, (f) summed up days of all heatwave during 2022. The definition of these metrics follows Hobday et al. (2016).

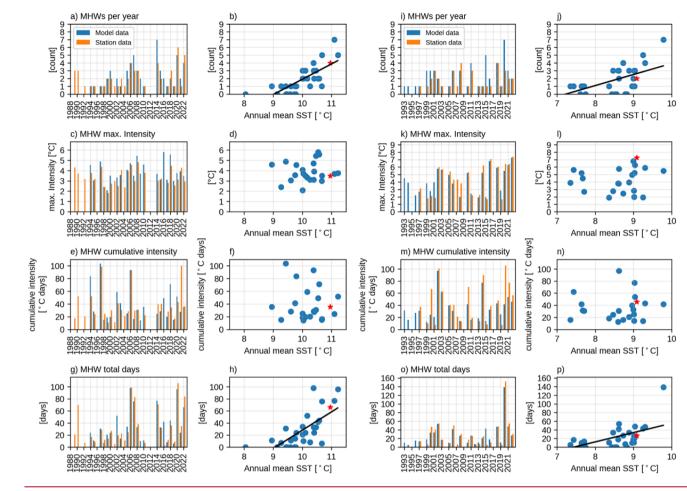
LT Kiel

Northern Baltic



LT Kiel

Northern Baltic

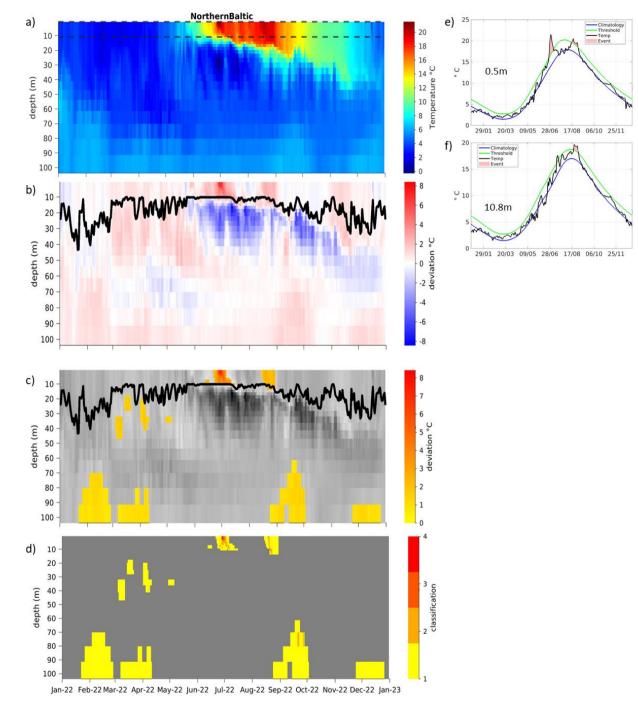


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Figure 24: Comparison and time series of annual MHW metrics (a,i: MHW events; c,k: maximum intensity [°C]; e,m: cumulative intensity [°C days]; g,o: MHW days) for station data (orange bars) and model data (blue bars) at the stations LT Kiel (left) and Northern Baltic (right). The MHW metrics from the model are plotted against the annual mean SST at that station with the year

Northern Baltic (right). The MHW metrics from the model are plotted against the annual mean SST at that station
 2022 marked in red. Statistically significant (95 %) correlations are indicated with a black line.

www marked in real statistically significant (25 70) correlations are indicated with a bla





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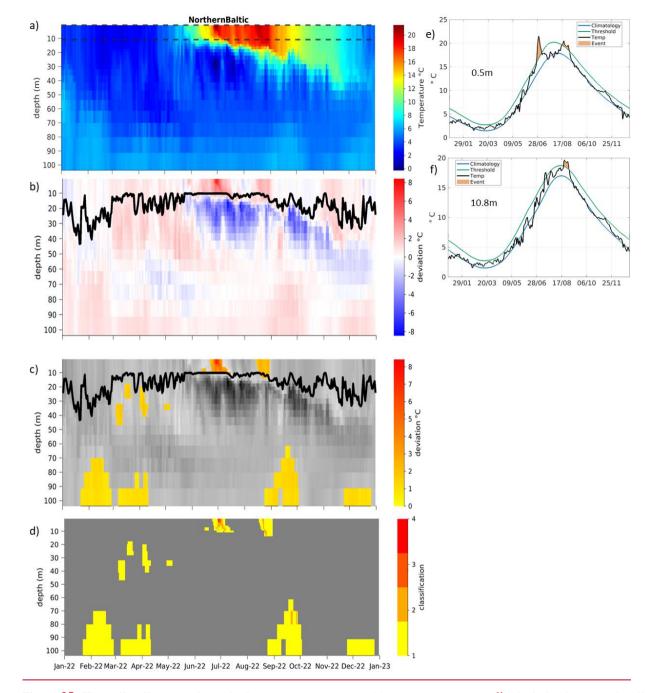
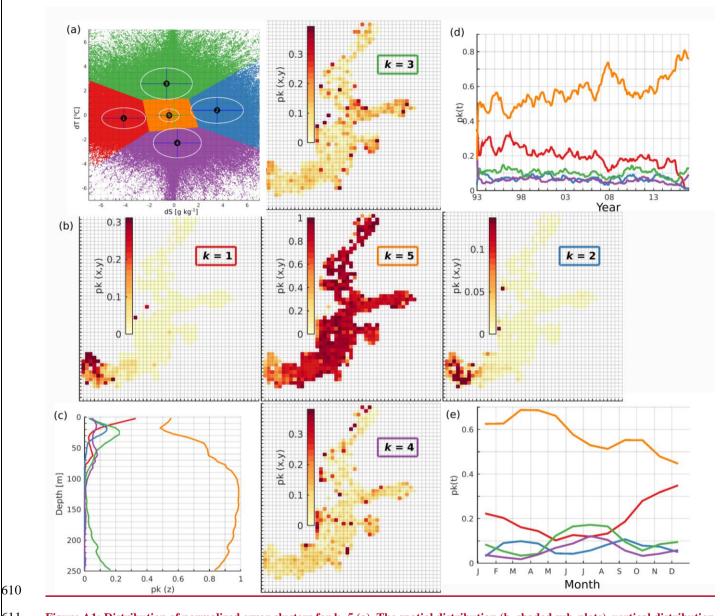


Figure <u>35</u>: Hovmöller diagrams show <u>absolute</u> water temperature (a), temperature <u>anomaliesdeviation between the climatology</u> and the <u>MYP data for 2022</u> (b) and MHWs (c) and their classifications (d, 1-moderate, 2-strong, 3-severe, 4-extreme) including the mixed layer depth <u>as the thick black line</u> (b and c) at Northern Baltic based on the Baltic Sea MYP (product ref. no. 3 in Table 1). The time series on the right (e-<u>hf</u>) are located at the vertical positions marked as dashed lines in (a) and show <u>SSTtemperature</u>

608 (black), climatology (blue), 90th percentile threshold for MHW analysis (green) and MHWs (red shading) based on model data at 609 depths of 0.5 m (e),) and 10.8 m (f), 80.1 m (g) and 103.8 m (h).). The period used for the climatology is 1993-2021.



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Figure A1: Distribution of normalized error clusters for k=5 (a). The spatial distribution (b, shaded sub-plots), vertical distribution (c), temporal distribution (d), and seasonal distribution (e) of the share of error points belonging to the five different clusters.

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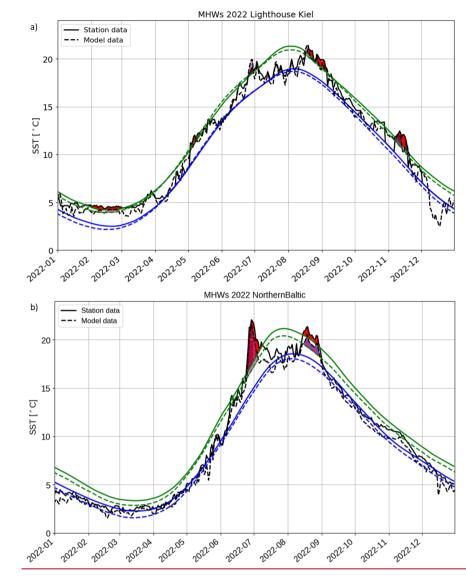


Figure A2: Comparison of station data with model data at (a) LT Kiel (product ref. no. 2 and 3 in Table 1), (b) Northern Baltic (K.
Hedi, FMI, pers. communication and product ref. no. 3 in Table 1). The dashed lines correspond to the model, while the continuous
lines correspond to the station data. In blue, the climatological mean is shown. The green lines show the 90th percentile threshold
for MHW detection and the black lines are the respective 2022 temperature data. The purple (model data) and red (station data)
marked areas show the detected MHWs in 2022. The reference period is 1993-2021 for LT Kiel (a) and 1997-2021 for Northern
Baltic (b).