

# 1 **Dissolved oxygen as indicator of multiple drivers of the marine** 2 **ecosystem: the Southern Adriatic Sea case study**

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7 **Abstract.** Oxygen is essential to all aerobic organisms and its dynamics in the ocean involve interconnected physical and  
8 biological processes that are at the basis for the marine ecosystem functioning. The study of dissolved oxygen (DO) variations  
9 under multiple drivers is currently one of the main goals of climate and marine ecological scientific communities, and the  
10 quantification of DO levels is essential for the assessment of the environmental status, especially in coastal areas.

11 We investigate the 1999-2021 interannual variability of DO in the southern Adriatic Sea, a marginal area of the Mediterranean  
12 Sea, where deep water formation processes occur, contributing significantly to the ventilation of the Eastern Mediterranean  
13 basin. Following the Marine Strategy Framework Directive, which promotes the integration of different observational  
14 platforms, we use DO modelled by the Copernicus Marine Mediterranean Sea biogeochemical reanalysis, which assimilates  
15 satellite chlorophyll concentrations and to which we apply a bias correction using DO Argo float measurements in 2014-2020.  
16 A correlation analysis of the time series of the first three modes of variability (86% of the total variance) of the DO profiles  
17 extracted from the bias-corrected reanalysis with key meteo-marine indicators shows a link with (i) net heat fluxes related to  
18 oxygen solubility, (ii) vertical mixing, (iii) biological production at the surface and in subsurface layers, and (iv) circulation  
19 associated with the entrance of northern Adriatic waters. The alternating entrance of Levantine and Atlantic Waters through  
20 the North Ionian Gyre (NIG) appears to be the driver of the fourth mode of variability, which explains 8% of the total variance.  
21 Moreover, we find that the first temporal mode of variability is the main driver of the negative anomaly of DO in the 0-600 m  
22 layer in 2021 with respect to the 1999-2020 climatology. We ascribe the lower content of DO in 2021 to a negative anomaly  
23 of the subsurface biological production in the same year, in agreement with the previous correlation analysis, but not to heat  
24 fluxes. Indeed, in agreement with previous studies, we observe a sharp increase in salinity favoured by the cyclonic circulation  
25 of NIG from 2019 onwards. We interpret this as a possible regime shift that is not captured by the time series analysis, and  
26 whose possible consequences for Ionian-Adriatic system ventilation and for marine organisms should be carefully monitored  
27 in the near future.

## 28 **1 Introduction**

29 Dissolved oxygen (DO) is a key indicator for monitoring the marine ecosystem functioning, because it is the result of several  
30 atmospheric, hydrodynamic and biogeochemical driving processes (such as air-sea fluxes, vertical convection and mixing,  
31 horizontal transport, biological production and consumption; Keeling and Garcia, 2002; Oschlies et al., 2018; Pitcher et al.,  
32 2021). Indeed, DO is currently being studied under the global warming scenarios by climate and marine ecological scientific  
33 communities (e.g. Pörtner et al., 2019; Kwiatkowski et al., 2020; Garcia-Soto et al., 2021), as oxygen depletion has been  
34 observed in the global ocean as well as at local scale (Breitburg et al., 2018). Climate models predict a reduction in global  
35 ocean dissolved oxygen content (Matear et al., 2000; Oschlies et al., 2008; Stramma et al., 2010, Reale et al., 2021), so this  
36 parameter is of primary interest especially in those areas where oceanic processes connect surface and deep layers.

37 The southern Adriatic Sea (SAdr, Fig. 1a) is one of these areas, as it is an area of deep water formation (Gačić et al., 2002;  
38 Pirro et al., 2022) and represents the deep engine of the eastern Mediterranean thermohaline circulation (Malanotte-Rizzoli et  
39 al., 1999), which is crucial for the eastern basin ventilation. The Adriatic Sea (Fig. 1a) is an elongated, semi enclosed and  
40 roughly north-south oriented basin characterized by a shallow northern shelf (shallower than 80 m) and a deep pit in its southern  
41 part (maximum depth of approximately 1200 m) which is connected to the Ionian Sea (central Mediterranean basin) through  
42 the Otranto Strait (with a maximum depth of 800 m). The Adriatic Sea is characterized by a cyclonic circulation governed by  
43 several drivers: river runoff, wind stress, surface buoyancy fluxes and mass exchanges through the Otranto Strait (Cushman-  
44 Roisin et al., 2013).

45 The SAdr is strongly influenced by the inflow of water masses from the northern Adriatic Sea (i.e., North Adriatic Dense  
46 Water, Querin et al., 2016) and the Ionian Sea. In particular, the inflow of southern water masses is triggered by the periodic  
47 reversal of Northern Ionian Gyre circulation (Gačić et al., 2002; Civitarese et al., 2010, Menna et al., 2019). This oscillating  
48 system, called the Adriatic - Ionian Bimodal Oscillating System (BiOS), changes the circulation of the Northern Ionian Gyre  
49 from cyclonic to anticyclonic and vice versa, modulating the advection of water masses in the Adriatic Sea (Gačić et al., 2010,  
50 Rubino et al., 2020). The cyclonic circulation of the Northern Ionian Gyre causes the advection of saline water masses of  
51 Levantine origin (i.e., Levantine Intermediate Water, Cretan Intermediate Water, Ionian Surface Water and Levantine Surface  
52 Water, Manca et al., 2006), while the anticyclonic circulation favours the inflow of Atlantic Water and a relative decrease of  
53 salinity in the SAdr (Gačić et al., 2011, Menna et al., 2022a). This feature has a strong influence on the biogeochemical  
54 properties of the SAdr, affecting nutrient availability (Civitarese et al., 2010), phytoplankton blooms (Gačić et al., 2002;  
55 Civitarese et al., 2010), and species composition (Batistić et al., 2014, Mauri et al., 2021).

56 While hydrodynamic and biogeochemical properties of SAdr have been widely described in several studies (e.g., Civitarese et  
57 al. 2010; Cushman-Roisin et al., 2013; Lipizer et al., 2014; Kokkini et al., 2018, 2019; Mavropoulou et al., 2020; Mihanović  
58 et al., 2021; Menna et al., 2022b), at the best of our knowledge DO dynamics in the area in connection with relevant driving  
59 processes over decadal time scales have not been addressed yet.

60 Investigating the DO multidecadal variability is crucial for quantifying the state of the marine environment (Marine Strategy  
61 Framework Directive, MSFD; Oesterwind et al., 2016) and for understanding anthropogenic impacts on the marine  
62 environment (Pörtner et al., 2022). The emerging ecosystem-based management method proposed by the MSFD (2008/56/EC)  
63 promotes the use of different observational platforms, allowing to synoptically collect information on the space-time  
64 distribution of important parameters related to water quality (Martellucci et al., 2021).  
65 In this context, the present work integrates the state-of-the-art approach of *in situ* measurements (in 2014-2020, distributed by  
66 Copernicus In Situ TAC) with the Copernicus biogeochemical reanalysis in the Mediterranean Sea at high resolution (Cossarini  
67 et al., 2021), with the aim of characterizing the DO dynamics in the SAdr in the 1999-2021 time period. In particular, we aim  
68 to assess DO inter-annual variability in an area (SAdr) sensitive to multiple drivers (e.g., atmospheric forcing, Mediterranean  
69 circulation, and biological processes) and to evaluate the relative importance of the different drivers in this area.

## 70 **2 Data and methods**

71 In the present study the DO concentration in the SAdr area (Fig. 1a) was assessed by combining data from the Copernicus  
72 reanalysis in the Mediterranean Sea (prod. ref. no. 1 in Table 1; Cossarini et al., 2021) in 1999-2021 and the Copernicus *in situ*  
73 dataset (prod. ref. no. 2, <https://doi.org/10.13155/75807>), available for the period 2014-2020 (Figs. 1b-c). The temporal  
74 evolution of the combined model-*in situ* DO concentration profile in 1999-2021 time period is shown in Fig. 1d.  
75 In particular, we used the BGC-Argo float measurements of *in situ* DO to compute a bias correction to the daily DO  
76 concentrations simulated by the biogeochemical reanalysis at  $1/24^\circ$  horizontal resolution. In fact, the biogeochemical  
77 reanalysis does not include BGC-Argo float DO assimilation and displays an average RMSD of  $15 \text{ mmol m}^{-3}$  for DO in the 0-  
78 600 m depth layer with respect to the observations in the area (Cossarini et al., 2021, Teruzzi et al., 2022). Quantile Mapping,  
79 a technique largely used for climate simulations (e.g., Hopson and Webster, 2010; Themeßl et al., 2011; Gudmundsson et al.,  
80 2012), was adopted to perform the reanalysis bias correction. The Quantile Mapping technique adjusts the cumulative  
81 distribution of the data simulated for the past or future period by applying a transformation between the quantiles of the  
82 simulated and observed data in the present. In our application, we adapted the code publicly provided by Beyer et al. (2020)  
83 at <https://doi.org/10.17605/OSF.IO/8AXW9> and included available *in situ* data of daily DO (Fig. 1c) within a representative  
84 area (Fig. 1b) of the southern Adriatic in the period 2014-2020, and DO reanalysis data for the same days of measurements.  
85 The representative area was identified by applying a spatial cross-correlation analysis (Martellucci et al., 2021) to the  
86 biogeochemical reanalysis centered on the SAdr pit and selecting the correlation threshold of 0.9 (Fig. 1b). Specifically, we  
87 considered the cross-correlation between the surface data of DO, nitrate and chlorophyll concentrations in the central point of  
88 the pit and those at each spatial grid point in the domain, to identify the area that displayed the same dynamics at the surface  
89 from a phenomenological perspective. Further details on the Quantile Mapping bias correction are included in Appendix A.  
90 We then applied the Empirical Orthogonal Function (EOF) analysis (e.g. Thomson and Emery, 2014) to the vertical profiles  
91 in Fig. 1d to describe DO variability in the SAdr area in the period 1999-2021. The EOF analysis allows us to identify the

92 spatial patterns of variability (i.e., EOF vertical patterns), describe how they change in time by means of time series (i.e., EOF  
93 time series), and associate the explained variance with each mode.

94 Finally, we performed a Pearson correlation analysis between the EOF time series in 1999-2021 and the following series of  
95 forcing indexes (reported in Fig. 2) providing evidence of the mechanisms driving oxygen concentration and dynamics in the  
96 area:

97 - heat fluxes in the SAdr as a proxy for thermal and mixing and stratification cycles (from prod. ref. no. 6 in Table 1; Fig. 2a);

98 - mixed layer depth in the SAdr as a proxy for both local vertical mixing and water residence times in the pit (prod. ref. no. 3;  
99 Fig. 2b);

100 - chlorophyll concentration at surface and in subsurface in the SAdr as a proxy for biological production in spring and late  
101 spring-summer, respectively (prod. ref. no. 1; Figs. 2c-d);

102 - heat fluxes in the northern Adriatic Sea (NAdr), as a proxy for dense water oxygen-rich formation in the NAdr and its  
103 transport into the pit (prod. ref. no. 6; Fig. 2e);

104 - Northern Ionian Gyre (NIG) vorticity derived from satellite altimetry, as a proxy of the inflow of Levantine waters and  
105 Atlantic Water (AW) (prod. ref. no 4 and 5; Fig. 2f).

106 In particular, the temporal phases of the NIG are defined as cyclonic and anticyclonic, respectively, when the vorticity field is  
107 positive and negative, as highlighted by the de-seasonalized time series in Fig. 2f.

108 Mixed layer depth (computed in prod. ref. no. 3 considering the  $0.03 \text{ kg m}^{-3}$  density difference with respect to the near-surface  
109 value at 10 m depth) and the chlorophyll at surface and in subsurface (30-80 m, where the deep chlorophyll maximum is  
110 located) were spatially averaged in the SAdr area ( $41.6^{\circ}$ - $42.1^{\circ}$ N;  $17.6^{\circ}$ - $18.1^{\circ}$ E, to consider the whole volume of the pit); heat  
111 fluxes were calculated in both the SAdr area and in the NAdr area ( $44.5^{\circ}$ - $45.5^{\circ}$ N;  $13^{\circ}$ - $13.5^{\circ}$ E), while current vorticity was  
112 computed in the Northern Ionian Sea ( $37^{\circ}$ - $39^{\circ}$ N;  $17^{\circ}$ - $19.5^{\circ}$ E).

113 In the correlation analysis, the time series of the heat fluxes in NAdr Sea (Fig. 2e) has been temporally lagged by 2 months, as  
114 an estimated mean time for the entrance in the SAdr pit of waters originated in the Northern Adriatic area (Vilibić et al., 2013;  
115 Querin et al., 2016; Mihanović et al 2021). Moreover, we tested the significance of the correlation coefficients between EOF  
116 and driver time series using a parametric t-test (with a reference significance level of 0.05).

## 117 **3 Results**

### 118 **3.1 Temporal scales of variability in connection with drivers**

119 Dissolved oxygen in the southern Adriatic area (Fig. 1a) shows in the subsurface layers an alternation between periods of  
120 enrichment (in 2004-2006, 2010-2013, 2016-2017) and sharp declines that impacted the Oxygen Minimum Layer (OML),  
121 located between 100 and 300 m. Low concentration values are observed also in the years between 1999 and 2003.

122 The EOF analysis was performed on the vertical profiles of the oxygen anomaly, derived by removing the mean profile in the  
123 period 1999-2021, and then normalized dividing by their standard deviation.

124 The time series of the first four EOF modes, which explain up to 95% of the oxygen variability in the water column are shown  
125 along with the corresponding vertical patterns in Figs. 3a,c,e,g and Figs. 3b,d,f,h, respectively.

126 The EOFs are interpreted considering the correlation of the EOF time series (Figs. 3a,c,e,g) with the time series of the forcing  
127 indicators shown in Fig. 2 (see Table 2), with heat fluxes in the northern Adriatic Sea time-lagged by two months as the  
128 estimated time of entry of the NAdr dense water in the SAdr pit.

129 The first mode (Figs. 3a-b), accounting for 48.9% of the explained variance, can be associated with the seasonal cycle of  
130 oxygen concentration in the upper layers: its vertical pattern mainly affects the first levels (Fig. 3b), the corresponding time  
131 series shows relative maximum values in spring (Fig. 3a) and it shows a statistically significant but moderate correlation  
132 ( $r=0.56$ ) with heat flux and a lower correlation with the subsurface chlorophyll concentration ( $r=0.43$ ) in the SAdr area (first  
133 column in Table 2).

134 The second and third modes (Figs. 3c-d and 3e-f, respectively), describing 19.7% and 17.7% of the variance respectively,  
135 affected both the upper and deeper layers. Both modes display relative maximum values in summer, but they have different  
136 correlation coefficients with the explanatory factors. The time series of the second mode (second column in Table 2) shows a  
137 significant but low correlation with multiple drivers, exceeding 0.4 only for surface chlorophyll ( $r=-0.41$ ) and waters from the  
138 NAdr area ( $r=0.48$ ). The time series of the third mode (third column, same Table) is moderately correlated with both surface  
139 chlorophyll ( $r=-0.61$ ) and NAdr waters ( $r=0.68$  correlation), but also with heat fluxes in the area ( $r=0.51$ ), and, to a lower  
140 extent, with mixed layer depth ( $r=-0.41$ ) and subsurface chlorophyll ( $r=0.48$ ).

141 The fourth mode (Figs. 3g-h), that describes 8% of the variance, can be ascribed mainly to the vorticity of the NIG ( $r=-0.37$ ,  
142 last column in Table 2), which affects the oxygen concentration in the intermediate layer (100-500 m depth), filled by LIW,  
143 and acts in the opposite direction in the upper and deeper layers (Fig. 3h).

144 Analysing the four modes in order of decreasing explained variance, we ascribe the seasonal variability connected with  
145 solubility mainly to the first mode, whereas we associate the biological contribution to oxygen dynamics to multiple interacting  
146 modes. In fact, the first mode explains the onset of the subsurface oxygen maximum (SOM) in spring, while the summer  
147 dynamics of the SOM is partially related to the third mode. The second mode, whose time series is correlated with surface  
148 chlorophyll evolution among other factors, can explain that part of the oxygen variability that is related to winter surface  
149 productivity.

150 The SOM, evident in summer oxygen profiles in Figs. 1c-d at about 40 m depth, is a feature that has already been observed in  
151 a great part of oligotrophic oceans (Riser and Johnson, 2008; Yasunaka et al., 2022) and of the Mediterranean Sea (e.g., Kress  
152 and Herut, 2001; Copin-Montégut and Bégovic, 2002; Manca et al., 2004; Cossarini et al., 2021; Di Biagio et al., 2022) and it  
153 represents an emerging property resulting from multiple interacting ecosystem processes (i.e., air-sea interactions, transport,  
154 mixing and biological production and consumption) and is, indeed, captured by multiple modes.

155 The third mode, which also describes the high concentration values in the deep layers in the period 2005-2006 and 2012-2014,  
156 is also moderately associated with a multi-annual signal of the inflow of deep denser and oxygenated water from the northern  
157 Adriatic Sea ( $r=0.68$ , third column in Table 2; Querin et al., 2016). Finally, it is worth noting that an EOF analysis of detrended

158 DO time series (not shown) yields fairly similar results but with the third mode only weakly correlated with the forcing indexes  
159 ( $r < 0.4$ ). Indeed, we can conclude that the third mode captures a signal of long-term evolution of oxygen concentration  
160 associated with changes in heat fluxes and chlorophyll concentration.

### 161 **3.2 The 2021 anomaly**

162 The year 2021 shows an overall negative anomaly in the oxygen concentration profile (Fig. 4b) compared to the 1999-2020  
163 climatological profiles (Fig. 4a). In particular, the anomaly affects a layer that thinned during the year, moving from 0-600 m  
164 depth in winter-early spring season to 30-400 m in late spring-summer and 0-80 m in fall. The (absolute) maximum values  
165 correspond to 25-30 mmol m<sup>-3</sup> at the surface in spring and at the SOM depth in summer.

166 We verified that, among the EOF modes, the negative anomaly of the first mode is the main contributor to the 2021 negative  
167 oxygen anomaly (not shown). The time series of the first mode (Fig. 3a) is actually negative from 2019 and corresponds to  
168 the negative anomaly of only one of its drivers (Table 2), i.e. subsurface chlorophyll (Fig. 2d), and not heat fluxes (Fig. 2a).  
169 In particular, we estimated a mean negative anomaly approximately equal to 6% with respect to the climatological mean (1999-  
170 2020) for subsurface chlorophyll in 2021.

171 One of the causes of the decrease in total oxygen concentration in the SAdr could be due to the exceptional salinization  
172 observed in the SAdr since 2017 (Mihanović et al., 2021, Menna et al., 2022b). This increase was related to the inflow of new,  
173 warmer and noticeably saltier water masses from the northeastern Ionian Sea (Mihanović et al., 2021, Menna et al., 2022b).  
174 The inflow of saltier and warmer water masses is also evident by observing the temporal evolution of these parameters through  
175 the Strait of Otranto (Fig. B1). In particular, in the upper layer (0-150 m) both temperature and salinity show an overall positive  
176 trend throughout the period 1999-2021, whereas the decrease observed in 2006-2011 and 2017-2018 can be associated with  
177 the inflow of less saline AW, triggered by the anticyclonic circulation of the NIG (Fig. 2f). In the intermediate layer (150-600  
178 m), salinity shows a positive trend in 1999-2021, while no clear trend is observed for temperature. Moreover, a sharp increase  
179 in salinity ( $\sim 0.1$ ) is observed in 2019. This increase occurred after the NIG inversion from anticyclonic to cyclonic (Fig. 2f),  
180 resulting in a further increase in salinity due to both the decrease in AW advection and the increase in LIW inflow.

### 181 **4 Conclusions**

182 Merging the Copernicus biogeochemical reanalysis in the Mediterranean Sea with *in situ* TAC data of biogeochemical Argo  
183 floats allowed us to characterize the interannual variability of dissolved oxygen in the southern Adriatic Sea in the 1999-2021  
184 time period and the 2021 anomaly with respect to the mean over 1999-2020. This study enriches our knowledge of the dissolved  
185 oxygen state and long-term dynamics in the area, by proposing a seamless time and space perspective that is complementary  
186 to previous climatologies and data aggregation information (e.g. Lipizer et al., 2014) and adding an explanatory framework  
187 for the driving mechanisms in the marine environment.

188 The EOF statistical analysis that we conducted on the vertical oxygen profiles yielded two key results. First, in contrast with  
189 a climatological view, the analysis was able to capture most of the inter-annual oxygen variability associated with variability  
190 of the main drivers (i.e., heat fluxes affecting solubility; biological productivity; vertical mixing). We do not detect a clear  
191 deoxygenation trend in the subsurface layer, while the multiannual variability is characterised by an alternation of enrichment  
192 and reduction phases, whose dominant correlations with the drivers for each EOF time series are in the (absolute) range 0.40-  
193 0.70. The possibility to observe such cyclic signals is enhanced by the relatively small volume and short residence time of the  
194 SAdr pit waters (Querin et al., 2016) with respect to other Mediterranean areas (Coppola et al., 2018). This feature makes the  
195 SAdr a potential efficient probe to detect a rapid response to changes in its meteo-marine drivers, i.e., circulation and  
196 atmospheric patterns.

197 Indeed, as our second result, the variability that is not explained by the EOF decomposition appears to be connected with a  
198 possible regime shift, associated with the entrance of new water masses, warmer, markedly saltier and less oxygenated, that  
199 were not previously observed in the analyzed time period.

200 The exceptional increase in salinity occurring after 2019 has been already documented (Mihanović et al., 2021; Menna et al.,  
201 2022b) and also observed north of the SAdr pit. Further monitoring of such anomalously high salinity values and assessment  
202 of their potential impact on the marine food web is of great importance, as picoplankton groups are sensitive to this  
203 environmental variable (Mella-Flores et al., 2011) and changes in biomass and production due to salinity have already been  
204 observed in previous studies in the Adriatic Sea (Beg Paklar et al., 2020; Mauri et al., 2021). Moreover, if such a strong  
205 negative oxygen anomaly as observed in 2021 were to persist, it could have direct impacts on local marine organisms, as well  
206 as on the cycling of dissolved chemical elements (Conley et al., 2009) potentially altering the energy flux towards the higher  
207 trophic levels (Ekau et al., 2010). The importance of the relationship between dissolved oxygen and the catch distribution of  
208 some marine species has already been proved in the Adriatic Sea (Chiarini et al., 2022).

209 By integrating model and *in situ* data, our study demonstrates the importance of following up the oxygen content in a seamless  
210 spatial and temporal way, as it is a fundamental indicator of good environmental status (GES, Oesterwind et al., 2016) and a  
211 factor that significantly affects fishing activities and economy.

212

## 213 **Appendix A: Quantile Mapping bias correction of DO concentration profiles**

214 Figures A1 and A2 show the modelled DO concentration profiles and histogram distributions before and after the Quantile  
215 Mapping bias correction, respectively, conducted by using the BGC-Argo float measurements available in 2014-2020 (Fig.  
216 1c). The Quantile Mapping, better than other methods (i.e. Additive Delta Change, Multiplicative Delta Change and Variance  
217 Scaling; results not shown), acted on the profiles by modifying the values of the concentrations (as indicated by the different  
218 colorbars in Figs. A1a and A1b) but, at the same time, maintaining the main dynamics observed before the correction: mixing  
219 and stratification at the surface during the year, subsurface oxygen maximum onset in spring and development in summer, and  
220 interannual variability related to the mixed layer depth dynamics in the intermediate layers. The distributions of the values of  
221 the model output before and after the Quantile Mapping bias correction and the values from BGC-Argo floats are displayed in

222 Fig. A2. The correction actually changed the modelled values (Fig. A2a) to reproduce the shape of the distribution of the  
223 observations (Fig. A2c). In particular, after the correction (Fig. A2b) the modelled data show higher variability and a more  
224 skewed distribution toward the higher values, similarly to the observations.

225

## 226 **Appendix B: Time series of surface and intermediate temperature and salinity at the Otranto Strait**

227

### 228 **Data availability**

229 Publicly available datasets were analyzed in this study. Modelling and *in situ* data can be found at the Copernicus Marine  
230 Service, with references and DOIs indicated in the Table 1 of the manuscript.

231

### 232 **Author contribution**

233 VDB and GC conceived the idea. VDB, RM and MM conducted the analysis. VDB, RM and GC wrote the first draft, with  
234 contributions from the other co-authors. All the authors discussed and reviewed the submitted manuscript.

235

### 236 **Competing interests**

237 The authors declare that they have no conflict of interest.

238

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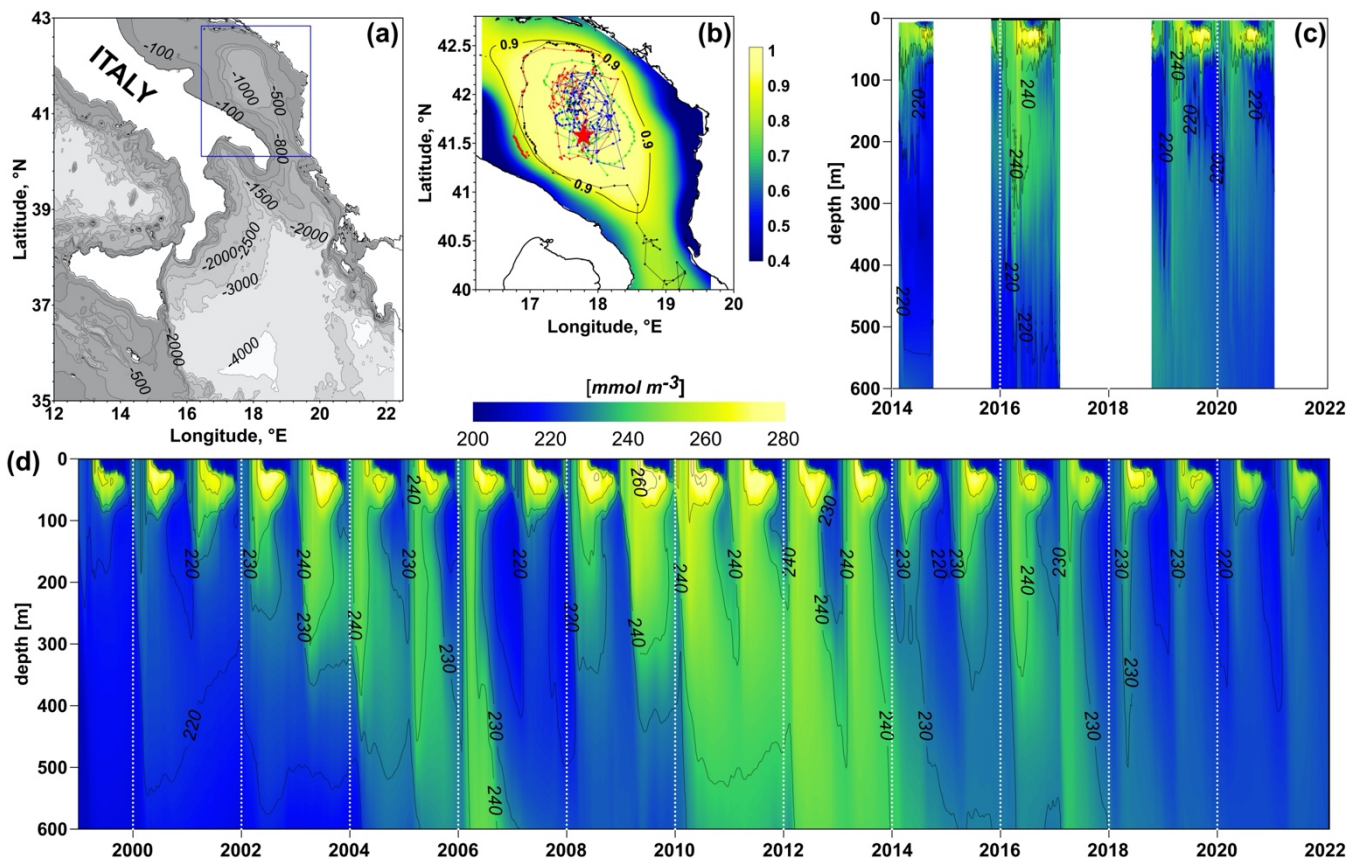
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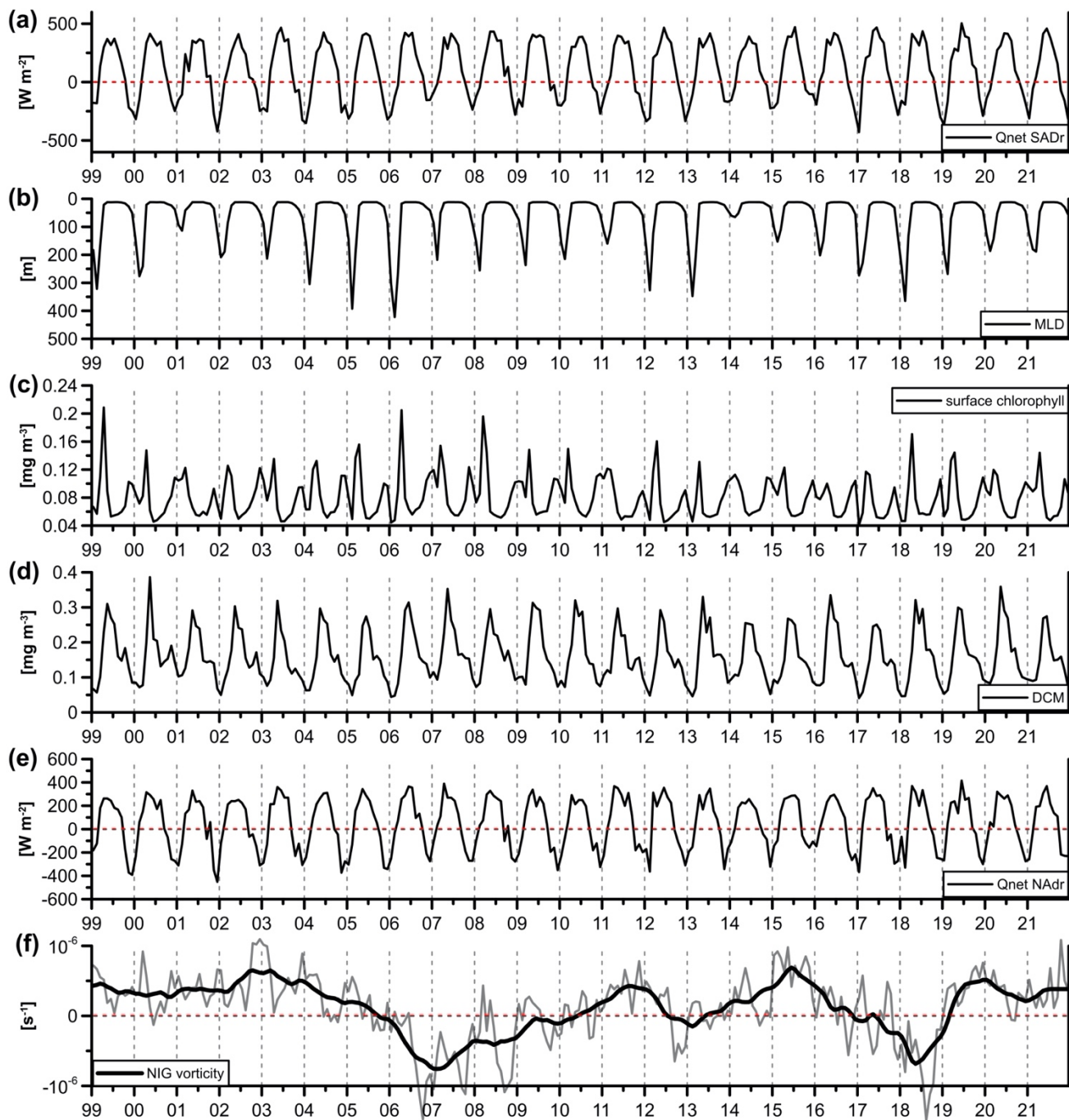
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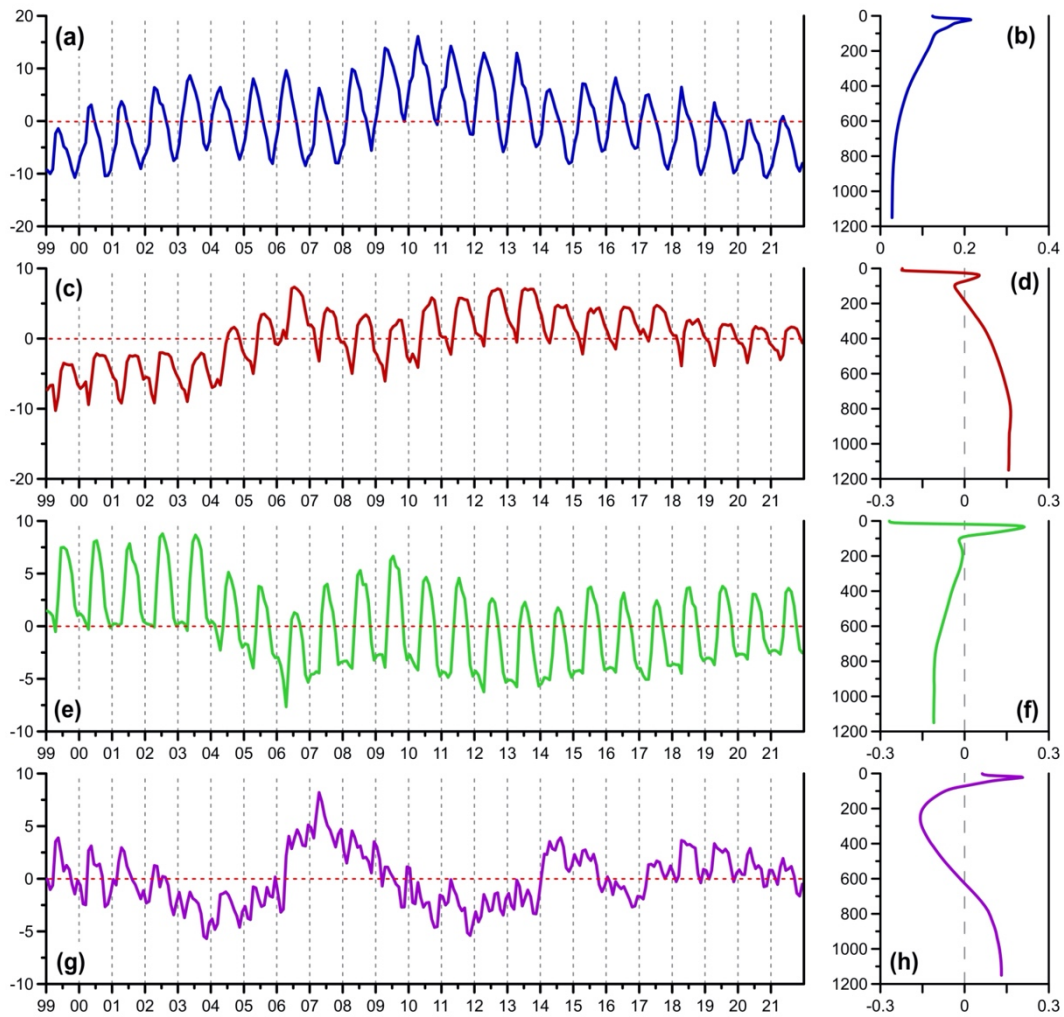
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516 **Figure 1:** (a) Southern Adriatic area (blue square) within Mediterranean Sea; (b) Cross-correlation map of surface oxygen, nitrate  
 517 and chlorophyll concentration provided by Copernicus biogeochemical reanalysis (prod. ref. no. 1, Table 1) in the southern Adriatic  
 518 area with respect to the central point of the pit indicated by the red star; the black contour line delimits the area with cross-  
 519 correlation equal or higher than 0.9; dashed lines indicate the trajectories of BGC-Argo floats (In Situ TAC data, prod. ref. no. 2)  
 520 passing the area. (c) Hovmöller diagrams of the dissolved oxygen concentration from In Situ TAC data (prod. ref. no. 2) within the  
 521 0.9 cross-correlation area (panel (b)). Data have been interpolated for readability of the plot. (d) Hovmöller diagrams of dissolved  
 522 oxygen concentration from Copernicus biogeochemical reanalysis (prod. ref. no. 1), spatially averaged within the area of cross-  
 523 correlation equal to 0.9 (panel (a)) in 1999-2021 time period, after the bias correction procedure based on In Situ TAC data (prod.  
 524 ref. no. 2).



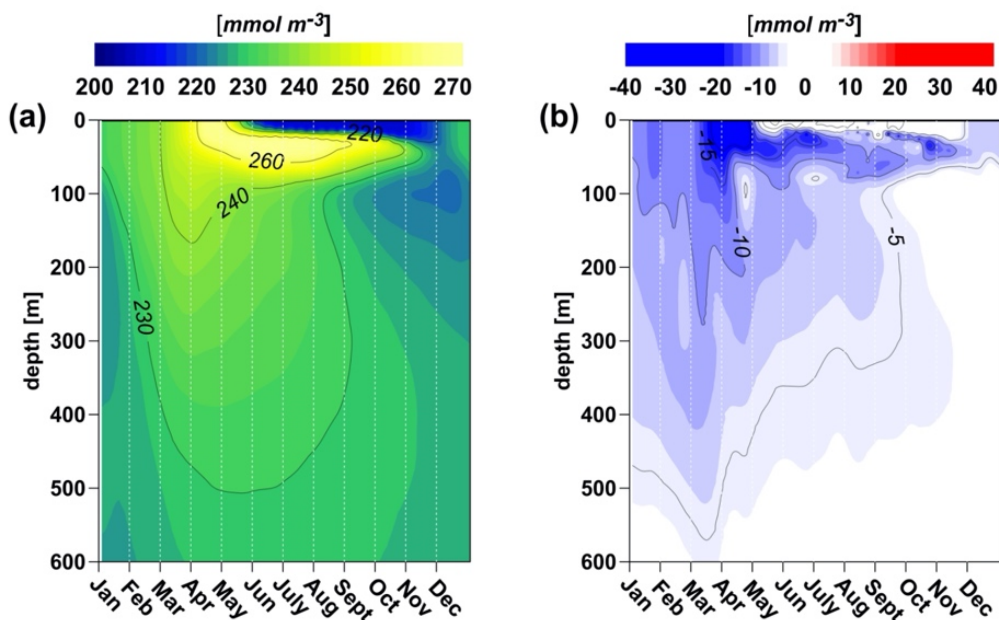
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526 **Figure 2: Time series of the main forcing in the 1999-2021 time period: (a) net heat fluxes in SADR (prod. ref. no. 6 in Table 1), (b)**  
 527 **mixed layer depth (prod. ref. no. 3), (c) surface chlorophyll concentration (prod. ref. no. 1), (d) subsurface chlorophyll concentration**  
 528 **(30-80 m layer in which deep chlorophyll maximum (DCM) is located, prod. ref. no. 1), (e) net heat fluxes in NADR (prod. ref. no. 6),**  
 529 **(f) NIG current vorticity (gray line) and de-seasonalized time series as obtained by applying a low-pass filter of 13 months (black**  
 530 **thick line) (prod. ref. no. 4 and 5).**



531

532 **Figure 3: EOF time series (a, c, e, g) and vertical patterns (b, d, f, h) of the first four modes computed on the bias-corrected dissolved**  
 533 **oxygen concentration in the southern Adriatic area shown in Fig. 1d. The explained variances of the four modes are: 48.9%, 19.7%,**  
 534 **17.7% and 8.4%.**



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537 Figure 4: Hovmöller diagrams of: mean over 1999-2020 of daily oxygen concentration computed from Copernicus biogeochemical  
 538 reanalysis (prod. ref. no. 1, Table 1) after the bias correction procedure based on In Situ TAC data (prod. ref. no. 2) (a) and anomaly  
 539 in 2021 with respect to the 1999-2020 period (b).

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Prod. ref. no.	Product ID & type	Data access	Documentation
1	MEDSEA_MULTIYEAR_BGC_006_008 Mediterranean Sea Biogeochemistry Reanalysis	<a href="https://doi.org/10.25423/CMCC/MEDSEA_MULTIYEAR_BGC_006_008_MEDBFM3">https://doi.org/10.25423/CMCC/MEDSEA_MULTIYEAR_BGC_006_008_MEDBFM3</a> (EU Copernicus Marine Service Product, 2022a)	Quality Information Document (QUID): Teruzzi et al., (2022) Product User Manual (PUM): Lecci et al., (2022a) Cossarini et al., (2021)
		<a href="https://doi.org/10.25423/CMCC/MEDSEA_MULTIYEAR_BGC_006_008_MEDBFM3I">https://doi.org/10.25423/CMCC/MEDSEA_MULTIYEAR_BGC_006_008_MEDBFM3I</a>	

		(EU Copernicus Marine Service Product, 2022b)	
2	INSITU_MED_PHYBGCWAV_DISCRETE_MYNRT_013_035 Mediterranean Sea-In-Situ Near Real Time Observations	<a href="https://doi.org/10.48670/moi-00044">https://doi.org/10.48670/moi-00044</a> (EU Copernicus Marine Service Product, 2022c)	Quality Information Document (QUID): Wehde et al., (2022) Product User Manual (PUM): In Situ TAC partners (2022)
3	MEDSEA_MULTIYEAR_PHY_006_004 Mediterranean Sea Physics reanalysis	<a href="https://doi.org/10.25423/CMCC/MEDSEA_MULTIYEAR_PHY_006_004_E3R1">https://doi.org/10.25423/CMCC/MEDSEA_MULTIYEAR_PHY_006_004_E3R1</a> (EU Copernicus Marine Service Product, 2022d)  <a href="https://doi.org/10.25423/CMCC/MEDSEA_MULTIYEAR_PHY_006_004_E3R1_I">https://doi.org/10.25423/CMCC/MEDSEA_MULTIYEAR_PHY_006_004_E3R1_I</a> (EU Copernicus Marine Service Product, 2022e)	Quality Information Document (QUID): Escudier al., (2022) Product User Manual (PUM): Lecci et al., (2022b) Escudier et al., (2021)
4	SEALEVEL_EUR_PHY_L4_MY_008_068 European Seas Gridded L 4 Sea Surface Heights And Derived Variables Reprocessed 1993 Ongoing	<a href="https://doi.org/10.48670/moi-00141">https://doi.org/10.48670/moi-00141</a> (EU Copernicus Marine Service Product, 2022f)	<u>Quality Information Document (QUID): Pujol al., (2022a)</u> <u>Product User Manual (PUM): Pujol et al., (2022b)</u>
5	SEALEVEL_EUR_PHY_L4_NRT_OBSERVATIONS_008_060	<a href="https://doi.org/10.48670/moi-00142">https://doi.org/10.48670/moi-00142</a>	Quality Information Document (QUID): Pujol al., (2022c)

	European Seas Gridded L 4 Sea Surface Heights And Derived Variables Nrt	(EU Copernicus Marine Service Product, 2022g)	Product User Manual (PUM): Pujol et al., (2022d)
6	ERA5 hourly data on single levels from 1940 to present Global climate and weather reanalysis	<a href="https://doi.org/10.24381/cds.adbb2d47">https://doi.org/10.24381/cds.adbb2d47</a> (EU Copernicus Climate Change Service Product, 2022)	Hersbach et al., (2018)

542 **Table 1: Products used in the present work. Prod. ref. no. 3 is a forcing for prod. ref. no. 1 and prod. ref. no. 6 is a forcing for prod.**  
543 **ref. no. 3. Complete references for prod. ref. no 1, 3 and 6 are reported in the bibliography.**

544

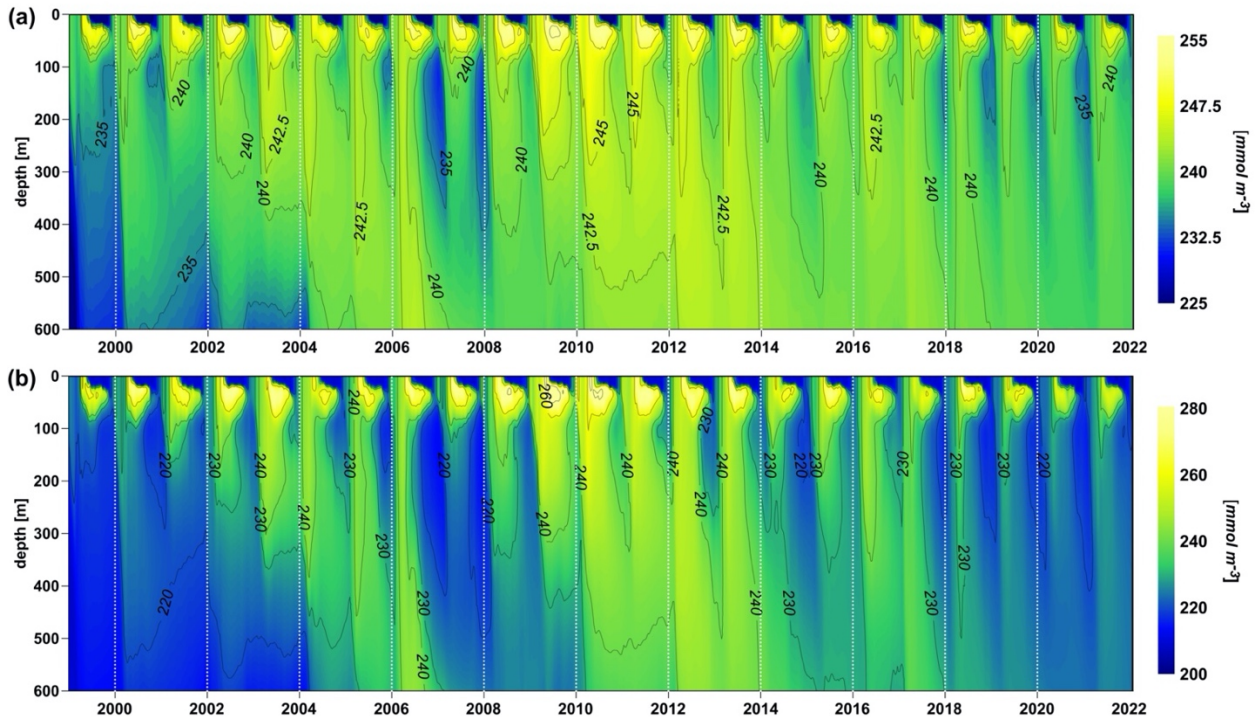
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	<b>mode 1</b>	<b>mode 2</b>	<b>mode 3</b>	<b>mode 4</b>
<b>HFlux (SAdr)</b>	0.56	0.15	0.51	0.32
<b>MLD (SAdr)</b>	n.s.	-0.28	-0.41	-0.25
<b>surf chl (SAdr)</b>	n.s.	-0.41	-0.61	n.s.
<b>subsurface chl (SAdr)</b>	0.43	0.13	0.48	0.34
<b>Hflux NAdr (2-months lagged)</b>	n.s.	0.48	0.68	0.16
<b>NIG vorticity (NIon)</b>	n.s.	-0.40	n.s.	-0.37

547 **Table 2: Correlations between the first four temporal modes of EOFs of DO (Figs. 3a,c,e,g) and the forcing fields (Fig. 2, with heat**  
548 **fluxes in the northern Adriatic Sea time-lagged by two months). Not statistically significant correlations are identified by a**  
549 **significance level higher than 0.05 and indicated by “n.s.” acronym in the table.**



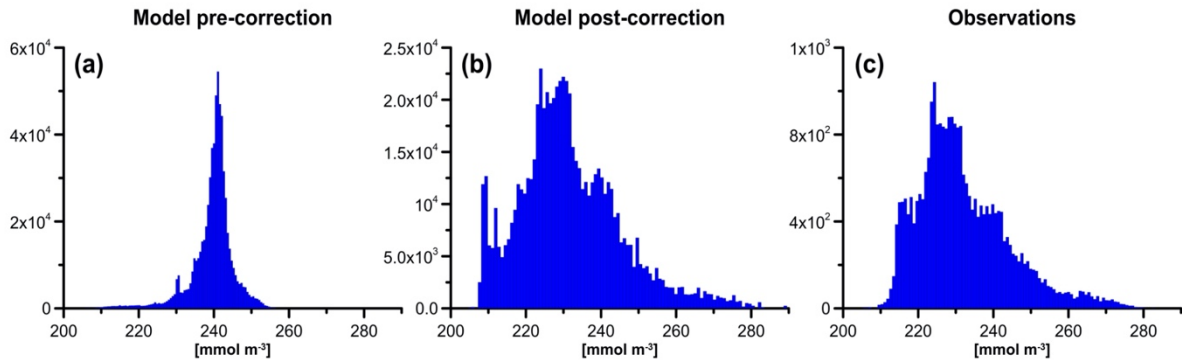


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551 **Figure A1: Hovmöller diagram of the modelled oxygen concentrations spatially averaged within the area of autocorrelation equal**  
 552 **to 0.9 indicated in Fig 1b, before the bias correction by Quantile Mapping (a) and after the procedure (b).**

553

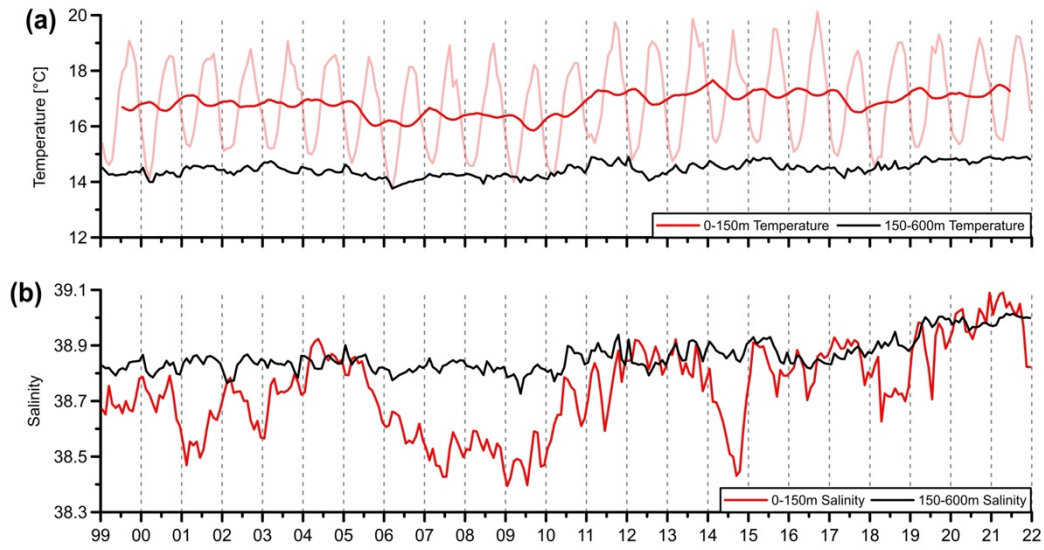
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556 **Figure A2: Frequency histogram of modelled oxygen concentrations before the bias correction by Quantile Mapping (a) and after**  
 557 **the procedure (b), compared with BGC-Argo observations (c).**

558



559

560 **Figure B1: Time series of temperature (a) and salinity (b), averaged in the vertical layers 0 - 150 m (red lines) and 150-600 m (black lines) of the Otranto Strait (39.8°N, 18.5° - 19.5° E) in the 1999-2021 time period. In the top panel, light red and dark red indicate**  
 561 **data before and after de-seasonalization, respectively. Data are provided by Copernicus physical reanalysis (prod. ref. no. 3, Table**  
 562 **1).**

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